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Escuela de Ciencias de la Tierra, Energía y Ambiente

TÍTULO: MAGNETOMETRY SURVEY APPLIED TO GEOTHERMAL EXPLORATION IN CHACHIMBIRO, NORTHERN OF ECUADOR

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obtención del título de Geólogo

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Dedication

To my beloved family, which always have been there to give me their unconditional love and support. Especially my grandparents Pepito and Catita, my parents Jhon and Marlene, and my brothers and sister David, Verónica and Pablo.

Javier Ricardo Pauta Ordóñez

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Resumen

Los estudios geofísicos son una forma eficiente de obtener información de las estructuras bajo la superficie. El método magnético es especialmente útil para detectar estructuras poco profundas y cambios en la magnetización debido a varios procesos, tales como el fallamiento y las alteraciones hidrotermales. A pesar de la riqueza de recursos geotérmicos en América del Sur, su uso para generación de energía eléctrica es muy limitado. En Ecuador, la mayoría de los proyectos geotérmicos están en fases de prospección y exploración inicial. Chachimbiro, ubicado al norte del Ecuador, es uno de los sitios con mayor potencial a convertirse en una planta de energía geotérmica. El objetivo de este trabajo es el de realizar estudios complementarios para mejorar el modelo existente del sistema geotérmico. Para esto, realizamos un estudio magnético de alta resolución con un distanciamiento de ~30 m entre puntos, alrededor del área del pozo Chachimbiro 1, y así comprender mejor las estructuras poco profundas que están sobre el reservorio, de la misma forma se realizaron dos líneas de estudio con un distanciamiento de ~5 m entre puntos y perpendicularmente a las fallas locales. Luego se compararon estos resultados con datos magnetelúricos y gravimétricos, que son menos precisos para detectar rasgos poco profundos. La comparación con estudios previos mostro que las fallas y la distribución de anomalías magnéticas de este estudio se ajustan con lo descrito en estudios previos. Nuestro estudio magnético fue útil para diferenciar las anomalías magnéticas relacionadas con topografía, fallamiento y alteraciones hidrotermales. Además, este estudio contribuirá a tener un mejor entendimiento del área, lo que es necesario para realizar futuras perforaciones en el área, que es potencialmente la primera planta de energía geotérmica del país.

Palabras Clave: Magnetometría, métodos Geofísicos, sistemas Geotermiales, volcán Chachimbiro, Detección de fallas.

Abstract

Geophysical surveys are efficient ways to obtain information on subsurface structures. The magnetic method is especially useful to detect shallow structures and changes in magnetization due to several processes, such as faulting and hydrothermal alterations. Despite the richness of available geothermal resources in South America, their use for electricity production is very limited. In Ecuador, geothermal projects in Ecuador are still at very initial phases of prospection and exploration. Chachimbiro, in northern Ecuador, is one of the potential site for developing a geothermal power plant. The objective of this project is to provide complementary studies to improve the existing model. We performed high resolution magnetometry survey with ~30 m spacing between points around the drilling area to get a better understanding of the shallow structure above the reservoir, as well as two survey lines with ~5 m spacing perpendicular to local faults. We then compared our results with existing magnetotelluric and gravimetric data, which are less precise for shallow structural features. Comparison with previous works shown that faults location and low magnetic anomalies distribution of this study fits with the locations described in previous studies. Our magnetic survey was usefulness to differentiate magnetic anomalies related to topography, faulting and hydrothermal alterations. Our study thus contributes to a better knowledge of the area, needed for future drilling planning of potentially, the first geothermal power plant of the country.

Keywords: Magnetometry, Geophysical methods, Geothermal systems, Chachimbiro volcano, Fault detection.

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1. Introduction

1.1. Fossil fuels and environmental impacts

Energy is a fundamental pillar in the development of the economy and technology. Nowadays, the principal sources of energy are fossil fuels, petroleum, coal, and natural gas. These sources are located under the ground and are a product of organic matter accumulation in anoxic conditions.

The main problem related to fossil fuel use is that their combustion releases several toxic gases (CO_2 , NO_x , SO_x , and CH_4 , also known as greenhouse gases) and other small particles (10 micrometers or less) to the atmosphere, causing environmental damage (Withagen, 1994; Williams, 2002). A significant accumulation of these gases leads to greater retention of the solar radiation that enters the Earth's atmosphere inducing global warming, what is known as the greenhouse effect. It is a natural process that would normally occur in periods of hundreds of thousands of years but is currently accelerating due to the increased use of these fuels (e.g. Al-Ghussain, 2019). Scientists have shown that this effect is linked to other environmental impacts such as melting of ice in the poles, rise of sea levels, acidification of oceans, among others (e.g. Karl and Trenberth, 2003). Moreover, fossil fuels are also considered finite resources, with estimates that they would run out by the middle to the end of the 21th century (e.g. Hoel, 1996; Dincer and Rosen, 1999). This concern has created the need to search for alternative energy sources that generate fewer pollutants to the environment, clean energies that can be exploited, generating significantly lower greenhouse gas emissions than fossil fuels. A good example of success in using alternative energy sources is Iceland, with almost 86% of its energy generation coming from renewable sources (68% geothermal, 18% hydropower). They use geothermal energy directly in industry, house heating, farming, and more (Ragnarsson, 2015).

Furthermore, these energies should be renewable, meaning that they will not be exhausted during human time scales. These new energies currently in use are the wind energy, hydraulic, solar, geothermal, and biomass. The largest inconvenience for the exploitation of renewable energies at a great scale is the high costs in comparison to fossil fuels. During the last decades, the prices have become more competitive, allowing the implementation of power plants that cover farms and houses' energetic needs, even for towns and cities. In some countries, renewable energies are the

cheapest alternative because of the abundance or prevalence of local energy sources (Dincer and Rosen, 1999).

Geothermal energy is the heat contained within the Earth. This energy, available almost everywhere on Earth's surface, can be recovered and exploited for human use. Geothermal heat mainly results from radioactive decay of potassium, thorium, and uranium in the crust (~50 %), with a small contribution from the friction along continental margins and remnant energy accumulated during primordial accretion. Earth is slowly cooling from its hot interior outwards to the cold atmosphere. The geothermal gradient is the average increase of the Earth's temperature with increasing depth. The average geothermal gradient of the Earth is about 2.5 – 3 °C every 100 m (Dickson and Fanelli, 2004). This average geothermal gradient is representative of most places on Earth while others, such as volcanoes, have higher gradients. The rocks' temperature increases much faster with depth at these locations, allowing hot water reservoirs to form at much more shallow depths.

1.2. Geothermal systems

Geothermal systems can be found in different parts of the world, but only those with a normal or abnormally high gradient are the most exploitable, generally related to tectonic margins. These systems are characterized by their temperature; low-temperature systems up to 100 °C and high-temperature systems range from 100 to >400 °C (Dickson and Fanelli, 2004). Three elements are essential to form a geothermal system: a heat source, a reservoir, and a fluid. The heat source can be magmatic intrusions that are relatively shallow (~10 km) or magmatic chambers at dormant volcanoes, from which the heat is conducted up to the overlying rocks. The reservoir is formed by a layer of porous rocks in which the fluids are accumulated. It is overlain by a layer of impermeable rocks (also called the cap rock) and connected to the surface by faults where water can be expelled. An efficient geothermal system requires replenishing the system's fluid with meteoric water (i.e., rainwater). Once heated, the geothermal water interacts with rocks increasing the number of cations, and can become acidic if the magmatic source releases magmatic gases. The mechanism that controls geothermal systems is fluid convection, occurring due to changes in fluid density. The heat from the source causes the thermal expansion of fluids, making them less dense. These fluids with low-density rise and are replaced by cooler and thereby denser fluids.

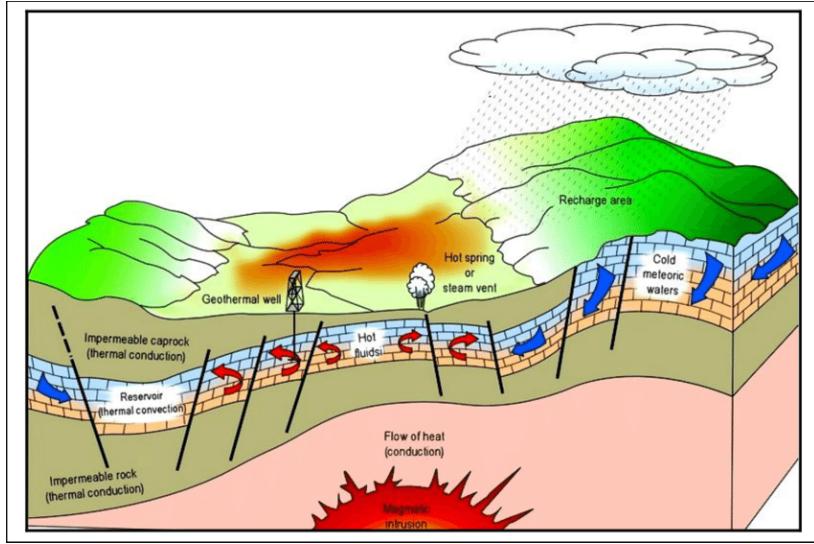


Figure 1: Schematic representation of an ideal geothermal system, image from Syukri et al., (2018).

The system can be explained as infiltrated meteoric water being heated by hot rocks and stored in a permeable reservoir. Convecting water transfers heat from the source to the surface, the reservoir being further refilled with meteoric water. An ideal geothermal system is represented in Figure 1 as a schematic model. Geothermal systems can be classified considering a variety of characteristics. Figure 2 shows a classification based on temperature, enthalpy, and physical state (Saemundsson et al., 2011).

Low-temperature (LT) systems with reservoir temperature at 1 km depth below 150°C. Often characterized by hot or boiling springs.	Lowenthalpy geothermal systems with reservoir fluid enthalpies less than 800 kJ/kg, corresponding to temperatures less than about 190°C.	Liquid-dominated geothermal reservoirs with the water temperature at, or below, the boiling point at the prevailing pressure and the water phase controls the pressure in the reservoir. Some steam may be present.
Medium-temperature (MT) systems with reservoir temperature at 1 km depth between 150- 200°C.		
High-temperature (HT) systems with reservoir temperature at 1 km depth above 200°C. Characterized by fumaroles, steam vents, mud pools and highly altered ground.	Highenthalpy geothermal systems with reservoir fluid enthalpies greater than 800 kJ/kg.	Two-phase geothermal reservoirs where steam and water co-exist and the temperature and pressure follow the boiling point curve. Vapour-dominated reservoirs where temperature is at, or above, boiling at the prevailing pressure and the steam phase controls the pressure in the reservoir. Some liquid water may be present.

Figure 2: Geothermal systems classification based on temperature, enthalpy, and physical state, according to Saemundsson et al. (2011).

Saemundsson (2011) also provide another classification based on the geological setting: 1) Volcanic geothermal systems (e.g geothermal field north of San Francisco, Surtsey, Iceland) in which the flow is mainly controlled by fractures and fault zones, associated with volcanic activity and located close to or inside volcanic complexes; 2) Convective fracture-controlled systems (e.g Soultz-sous-Forêts in France) with a heat source from crust at depth in tectonically active areas. Geothermal water flows at depths more than 1 km, mainly through vertical fractures; 3) Sedimentary geothermal systems (e.g. geothermal systems found along United States pacific coast, such Appalachia) present in several major sedimentary basins worldwide in zones with an anomalous geothermal gradient. This type of system is primarily conductive, but some convective systems may be contained in sedimentary rocks; 4) Geo-pressured systems related to oil and gas reservoirs (e.g. Texas and Luisiana gulf coasts) where fluids caught in stratigraphic traps have lithographic pressure; 5) Hot dry rock (HDR) or engineered geothermal systems (EGS) (e.g. some systems in Japan, Australia, France, Germany), consisting of the volume of rock heated by high heat flow or volcanism with relatively low or no permeability; 6) Shallow resources, referring to the usual heat flux in near-surface formations and thermal energy that is stored in rocks and warm groundwater systems near the surface.

This variety of geothermal systems implies that the Earth's available thermal energy is immense, but people can use only a modest fraction. Human use is limited to areas where geological conditions permit a carrier (water in the liquid phase or steam) to transfer the heat from deep hot zones to or near the surface (Lund and Boyd, 2016). The amount of usable energy from geothermal sources varies with depth and by extraction method. Figure 3 shows some of the worldwide uses of geothermal energy.

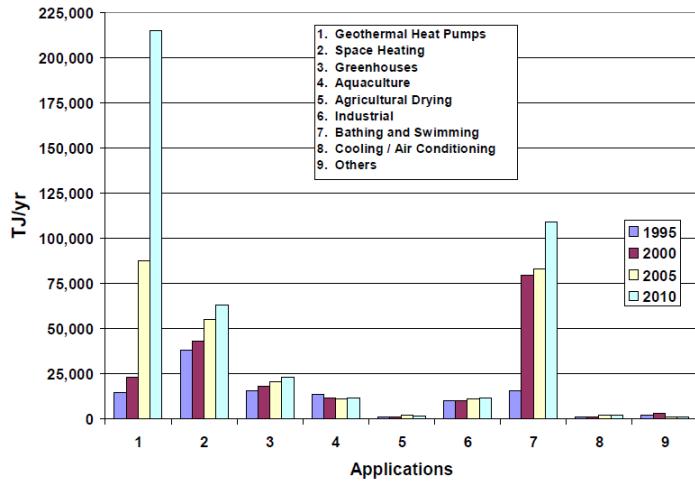


Figure 3: Comparison of worldwide geothermal energy utilization in TJ/year for 1995, 2000, 2005, and 2010. Image from Lund et al., (2010).

1.3. Geothermal prospection

The primary purpose of geothermal exploration is to define the reservoir's location, volume, shape, and structure and determine its characteristics such as fluid type, temperature, and how much energy can be produced. Exploration can be performed using geological, geochemical, and geophysical data. Geological prospection is usually the first step in geothermal prospection. It starts with observations of surface features like fumaroles, geysers, hot springs or steaming ground, or any other feature that can give information about the system and map faults and hydrothermal alteration regions (Gupta and Roy, 2006).

1.3.1. Geochemical prospection

Geochemical prospection relies on the chemical analysis of liquids and gases from a geothermal system to determine the nature and temperature of fluids in the reservoir. According to Óskarsson and Ármannsson (2015), geothermometry (fluid temperature estimate based on chemical composition) is based on the following assumptions: 1) The fluid reached the local equilibrium with secondary minerals in the reservoir at a specific temperature. The chemical equilibria have to be sensitive to temperature. 2) The composition of the fluids does not change during their rise to the surface. Solute geothermometry is based on minerals' solubility in the water and is measured by the concentration or activity, commonly utilizing silica geothermometers for it. 3) Finally, gas geothermometry assumes the gases equilibrium and uses gas geothermometers to estimate

temperatures from gas concentration in steam. The most common gases used are the geothermal gases CO₂, H₂S, H₂, and CH₄. Phuong et al. (2012) used geothermal fluids applied in geothermal prospection in Indonesia. They did a soil gas survey for radon (Rn), thoron (Tn), CO₂, mercury (Hg), and the chemical analysis of the hot spring waters in the Ungaran geothermal field. Chemical analysis showed a water temperature range from ~18 °C to ~56 °C, while the gas soil survey revealed several fault systems trending NNE-SSW and WNW-ESE. The presence of these same gases (Rn, Tn and CO₂) also helped identify a fracture zone that is allowing the fluids to migrate to the surface (Phuong et al., 2012).

1.3.2. Geophysical prospection

Geophysical surveys are the only way to obtain a detailed delineation of the subsurface structures, other than drilling (Gupta and Roy, 2006). These methods take into consideration the physical properties of the Earth. A geophysical survey may be helpful to delineate the geothermal area, and locate aquifers and other geological structures. Some of the physical factors with importance in a geothermal system are temperature, porosity, permeability, fluid salinity, and pressure (Georgsson, 2009). Direct geophysical methods applied to the Earth's surface cannot measure these properties. However, some methods may give indirect information about the geothermal systems: electrical conductivity, propagation of seismic waves, density, magnetic susceptibility (Dickson and Fanelli, 2004). According to Georgsson, one should distinguish between direct and indirect methods. Direct methods provide information on the factors affected by geothermal activity and include thermal, electromagnetic, electrical methods, and self-potential. On the other hand, passive methods give information that may reveal structures and geological bodies that would help understand the system, including magnetic surveys, gravity surveys, active seismic methods, and passive seismic monitoring.

Thermal methods consist of taking direct measurements in the field to determine temperature. These methods would help to determine the Earth's surface temperature, thermal inertia and thermal radiation from Earth's surface. It is limited to shallow depths (~10 m). Drilling is an active method that can be combined with measurements of the ground temperature, but it is usually expensive (Domra Kana et al., 2015).

The transient electromagnetic method (TEM) creates a grounded dipole that emits a constant magnetic field. A secondary induced field decays through time and it is monitored by measuring the voltage on the surface. TEM measurements allow the study of structures at depths from 1-1.5 km. Studies in Iceland helped to create maps down to 600 m depth, showing a low-resistivity-body with a high-resistivity core that reaches temperatures above 250 °C (Georgsson, 2009). Magnetotelluric method (MT) measures the Earth's natural electric field and magnetic field in orthogonal directions, both being dependent on the subsurface resistivity (Sircar et al., 2015). MT detects resistivity anomalies associated with faults, or the presence of a cap rock (i.e. low conductivity materials) and is thus useful to deduce the subsurface geology. MT studies from Sircar et al., (2015) in Arabia Saudi helped create 2D and 3D models which show the structure of the reservoir, composed by a shale/sandstone body between two layers of high-resistivity basalts. Electrical methods the principle says that the distribution of electrical potential in the ground around a current-carrying electrode depends on the electrical resistivities and distribution of the surrounding soils and rocks. This is useful because the electrical properties of the rocks are affected also by for example temperature and water content (Manzella, 1999). Electrical surveys show anomalies with low resistivity associated to a subsurface reservoir. The self-potential (SP) method is useful to understand the groundwater movement in the system.

The seismic method uses seismic waves that travel along the surface and subsurface through different rock material, and that are refracted or reflected by discontinuities. This allows a user to measure the velocity distribution, anomalies and attenuation of seismic waves. The method can be employed in a passive or active way (Domra Kana et al., 2015). Active methods use hammers or explosions to create seismic waves, and the information that can be obtained are the density of the formations, porosity, texture, boundaries and discontinuities. Passive methods use natural seismic activity and are useful to delineate active faults and permeable zones, as well as predict locations of hot bodies.

The gravimetric method consists of measurements of gravity anomalies in the Earth's surface derived from the density differences of the rocks in the subsurface. This method allows to determine masses of rock with high or low density which is possible by comparing it with the gravitational field measured (Georgsson, 2009). It also allows to search small, local, geological

structures (Hammer, 1939). Studies from Hochstein and Hunt, (1970) in New Zealand helped to provide the approximate depth of the greywacke basement beneath the Broadlands area.

Magnetic methods measure the Earth's magnetic field intensity, usually vertical magnetic gradient or total magnetic field. Anomalies in the magnetic field are due to differences in magnetic susceptibilities (i.e. differences in magnetizations) and they are often produced by remanent magnetism, carried by a ferrous body. The magnetic method is useful to locate intrusive bodies, tracing individual buried dykes and faults or estimating their depth (Bjornsson, 1980). Common sources of anomalies include dykes, faults, lava flows or iron-rich sediments (Chandra, 2015). Magnetic anomalies are useful to delineate high-temperature hydrothermal/geothermal systems; this is possible because several geothermal processes may alter the rock magnetization, going through a demagnetization of minerals or altering them to a less magnetic mineral (Caratori Tontini et al., 2016). Magnetic studies in Mahallat, Iran shows a 3D model of the area with an igneous body at a depth of approximately 1 km (Mohammadzadeh-Moghaddam et al., 2012).

1.4. State of geothermal prospection in Ecuador

To this day, geothermal energy in Ecuador has been exploited for direct uses only, such as bathing resorts and swimming pools, despite the fact that geothermal exploration started in 1978 under the supervision of the former Instituto Nacional de Electrification (INECEL) with the help of several specialists in geology, geochemistry and geophysics. This group explored several areas with recent volcanic activity or with surficial features like hot springs. In 1980 they determined priority geothermal areas, based on location, volcanic features (intensity, volume, frequency and age), hydrological conditions and surficial chemical characteristics. Following this, a collaborative study with Colombia to determine the viability of the Tufiño-Chiles-Cerro Negro area was initiated in 1983, and several geophysical studies were done in 1989 in Chachimbiro, in the northern part of Ecuador. All geothermal projects in the country ended in 1993 and were abandoned until 2007 (Piedra Lara, 2011). In 2008, there was a new proposal to change the productive and energetic matrix of the country, leading to studies in geothermal areas again. From 2010 until today, the Corporación Eléctrica del Ecuador (CELEC EP) has carried out studies in several geothermal areas of the country (Villacreses Baque et al., 2017).

In the last years, government policies have aimed for the development of renewable resources such as wind, solar, biomass and geothermal to reduce the use of fossil fuels and related gases emission. During the last decade, a geothermal plan was launched for electricity generation based on 11 potential geothermal sites: Chachimbiro, Chalpatán, Chacana-Jamanco, Chalupas, Guapán, Chacana-Cachiyacu, Tufiño, Chimborazo, Chacana-Oyacachi, Baños de Cuenca and Salcedo (Beate and Urquiza, 2015). Geological, geophysical and geochemical measurements have been done for Chachimbiro, Chacana-Jamanco and Chacana-Oyacachi projects (Beate and Urquiza, 2015). However, the only progressing project to date is that of Chachimbiro, with the first prospection well drilled in 2017 and more planned for the coming years.

2. Problem statement

Ecuador is a country with high volcanic activity. Previous government studies with assistance of foreign programs determined a geothermal potential at 1700 Megawatts electric (MWe) in 1999. Despite this great potential, there are still no geothermal power plants in Ecuador, nor an extensive use of this resource. One of the reasons, besides the investment problem, could be the lack of knowledge on how to use various tools for geothermal exploration and how to interpret the data. In Chachimbiro, the geothermal potential is estimated to be 81 MWe (Lloret and Labus, 2014). Several geophysical and geochemical surveys have been done, starting with a pre-drilling stage in 2012 and a drill stage in 2017. The methods used in that stage were magnetotellurics, gravimetry and magnetometry. However, many of these studies were led by foreign entities, without creating the local capacity to continue or expand these prospection surveys. The present study aims at being a complementary study of the main focus area in Chachimbiro, utilizing magnetometry to try to determine hydrothermal alteration zones, cap rocks and fault locations in the drilling area.

3. Objectives

- Collect the available geophysical and geological information.
- Perform a large-scale ground magnetometry survey with an evenly spaced grid in the area of the drilling sites.
- Analyze the magnetometry data to add knowledge about the subsurface geology of the area.
- Compare it with the newly collected data to understand the area's shallow geology.
- Determine the usefulness of magnetometry to locate faults in geothermally altered terrain.

4. Chachimbiro volcanic complex

4.1. Geological setting

The interaction of the Nazca, the South American and Cocos Plates is the main cause of the volcanism and tectonism in Ecuador. This resulted in the formation of the Andean Cordillera, which is conformed of two mountain ranges, the Cordillera Occidental to the west and the Cordillera Real to the east, with the Inter-Andean Valley depression in between (Figure 4A).

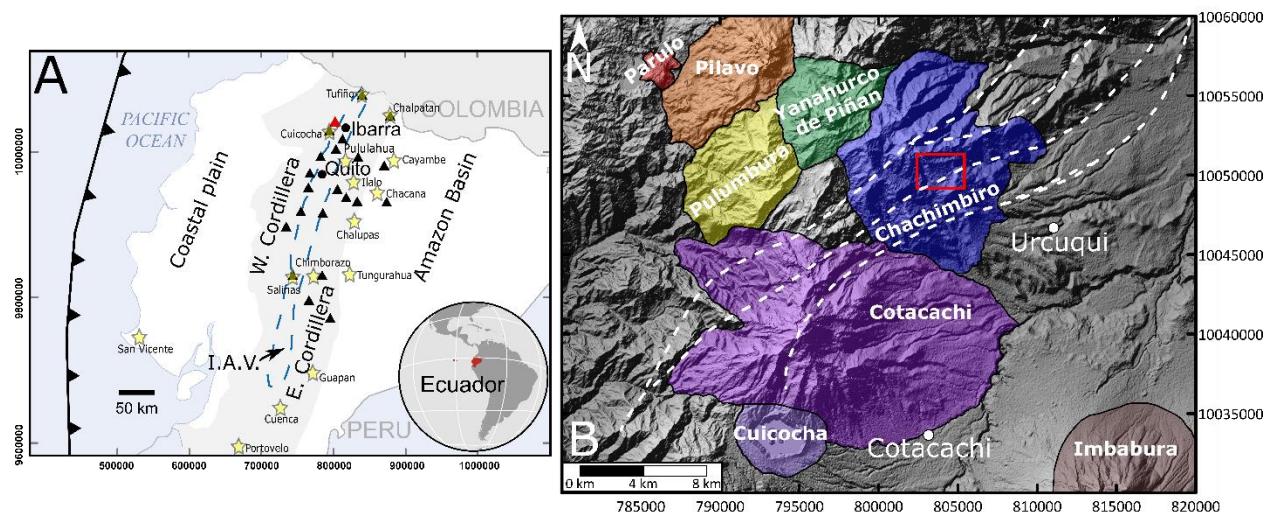


Figure 4: Location map. A) geologic map of Ecuador with the location of the study area in red square; black triangles are Quaternary volcanoes while red triangle is the Chachimbiro volcano. Yellow stars show the location of other geothermal areas in Ecuador. B) Geologic map of the area with the principal volcanic complexes close to Chachimbiro; white dots show Urcuqui and Cotacachi towns and dashed lines represent major faults in the zone. Red square represents the study area.

The Chachimbiro volcanic complex is located in Northern Ecuador within the Imbabura province (Figure 4B), marking the limit between the Western Andean Range and the Inter-Andean Valley. The average elevation around the Chachimbiro complex is 2560 m, and the topography is dominated by the Cotacachi and Yanaurco de Piñan stratovolcanoes. Chachimbiro, alongside Cotacachi, Cuicocha, Pulumbura, Yanaurco de Piñan and Pilavo volcanoes, belongs to the Andean Volcanic front of Ecuador. Except for Cuicocha, all the volcanoes in this zone are considered dormant volcanoes.

The volcanic edifice of Chachimbiro has a diameter of ~12 km, rising 1500 m above the Inter-Andean Valley. The basement of the Chachimbiro complex comprises Cretaceous rocks accreted

in a subduction zone. To the west of the complex, the basement is composed mainly of basic and intermediate volcanic rocks, emplaced in a submarine environment and covered by discontinuous turbidite deposits, pyroclastic and epiclastic continental deposits that belong to Silante formation (Granda, 2011), deposited during the Late Cretaceous to Early Miocene (Vallejo Cruz, 2007). To the east, the basement is composed of distal volcanic deposits covering sedimentary deposits from Miocene and Pliocene, which are affected by active NNE and NE faults with sinistral movement (Bernard et al., 2009), which determine the deposition zone during eruptive events. These faults are also zones of high degassing (Bellver-Baca et al., 2019). The Pallatanga formation, mainly composed of basaltic rocks, is considered the volcanic basement of the Western Cordillera. The Rio Cala Group overlays the Pallatanga formation and is mainly composed of massive basaltic to andesitic lavas, volcanic breccias, and volcanoclastic sandstones. The Natividad formation is related to the Rio Cala Group, composed of sedimentary rocks of the complex, correlated with turbidites deposited during the Eocene. This formation is exposed on the southeastern flank of the complex (Granda, 2011).

The eruptive history of Chachimbiro is described by Bernard et al., 2009. According to this study, the eruptive history is divided into four periods; three of them formed the volcano's domes. The first activity started forming the Huanguillaro dome during the Middle Pleistocene, mainly composed of calc-alkaline andesitic lava. The dome subsequently was destroyed by a giant landslide, depositing debris-avalanche deposits up to 25 km away from the source and leaving a caldera at the edges of the Huanguillaro and Conrayaro. The Tumbatu edifice formation characterizes the second period, growing within the depression of the collapsed Huanguillaro during the Late Pleistocene. Three main stages are responsible for the formation of this edifice; the first stage is characterized by the extrusion of dacitic domes producing large pyroclastic flows, resulting in block and ash flow deposits. More acidic lavas and explosive activity characterize the second stage, dated approximately 44 kyr. The corresponding units are ash and pumice flow deposits, reaching a thickness of 70 cm and traveling as far as 25 km from the source. The last stage is characterized by Plinian-type eruptions followed by new dacitic dome extrusions with block and ash flows and later emissions of andesitic tephra. This edifice also suffered a landslide related to an explosive event, forming a debris-avalanche deposit that even reached the Chota valley. The third period is characterized by the extrusion of several dacitic domes. One of the major eruptions resulted in the 3640 – 3510 BC blast, during which a dome located at the foot of La

Viuda peak violently exploded in a blast directed towards the south (Bernard et al., 2014). This extruded rhyodacite dome was formed by magma originating from two magmatic reservoirs emplaced at ~14.4 and 8 km depth, with temperatures of ~940 °C and ~860 °C, respectively. The last activity was the formation the Huga dome, accompanied by emission of andesitic tephras. Then, the volcano had a quiet period with high erosion in the volcanic edifice and a later remobilization of volcanic slope during a final dacitic tephra emission stage associated with Loma Albuji.

Previous studies observed several hydrothermal alterations which are: mesothermal propylitic (chlorite-epidote-calcite); which affected basaltic rocks not related with current thermal activity; Epithermal propylitic (smectite-chlorite), related with acid fluids and current thermal activity; argillic (smectite-kaolinite), related with hot springs and fumaroles in the area, also considered to be related with the current thermal activity; Advanced argillic (opal-smectite-kaolinite), which is related with fumaroles with high H₂S concentrations, and also considered related with the current argillic alteration; Carbonization, related with carbonate minerals deposited along structures where CO₂ gas is released (Pilicita, 2016).

4.2. Chachimbiro geothermal system

Previous works, such as the magnetotelluric (MT) studies from Pilicita (2016), determined that hydrothermal alteration is controlled by the right lateral strike - slip Azufral fault system (which includes the Pijumbi fault), trending NE-SW, and the Chachimbiro fault system and NW-SE lineaments. CO₂ and H₂S emissions are found close to the areas with hydrothermal alterations. Gas flow and alteration disappear in the NW direction from this point and no hot springs exist further than the Pijumbi and Azufral stream intersection. This may suggest that the Pijumbi fault is the boundary for the geothermal system. In terms of fluid geochemistry, studies showed springs with mixed chloride-bicarbonate composition and maximum temperature reaching 61 °C, indicating reservoir temperature of between ~235 and ~265 °C at the deepest part of the reservoir (Aguilera et al., 2005; Inguaggiato et al., 2010; Gherardi and Spycher, 2014). The main stress of the regional structural system is oriented in an E-W direction, creating a subparallel fault system oriented NE-SW, including two contemporaneous structures, the Azufral fault and Chachimbiro fault systems. This would allow the upward movement of fluids in the system.

Also, Bernard et al. in 2014 studied the last volcanic eruption between 3640 – 3510 years BC resulting in an extruded rhyodicite dome in between ~650 m wide and ~225 m high. Two reservoirs located 6.3 km east from the magma chamber were emplaced at ~14.4 and 8 km depth. The temperatures estimated were ~940 °C and ~860 °C. The depth of the clay minerals is described by Guillén, (2020), which estimate a smectite layer around ~200 m and an illite layer around ~250 m depth.

4.2.1. Previous geophysical studies

Several geophysical studies were carried out in the geothermal area of Chachimbiro, including Magnetotellurics (MT), Transient Electromagnetic Method (TEM), and resistivity method. Those studies helped to have an idea of the geothermal system of Chachimbiro, the estimated depth of the cap rock, and the degree of hydrothermal alteration in the geothermal area of Chachimbiro.

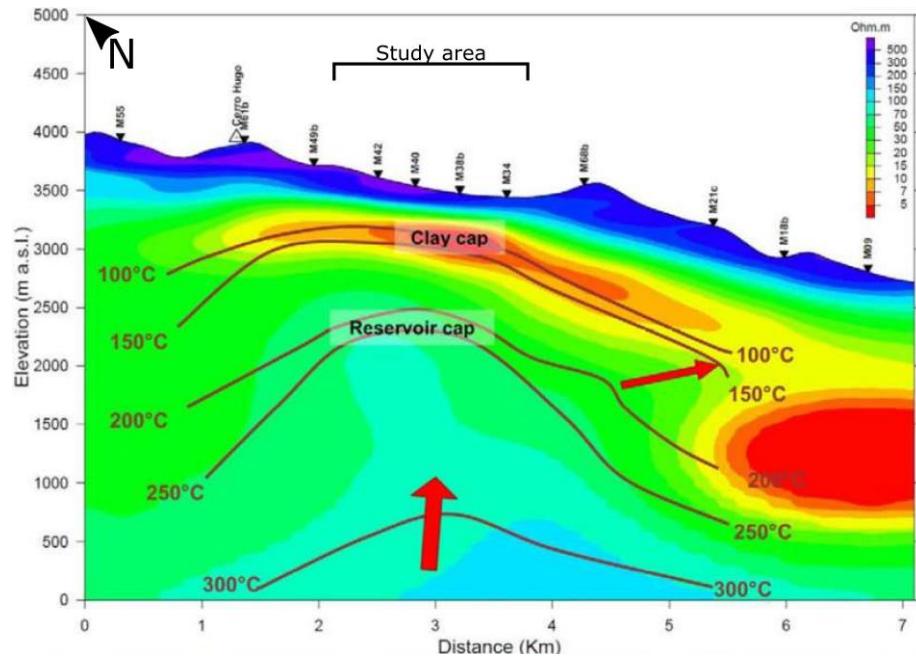


Figure 5: NW-SE resistivity cross-section, modified after Torres Calderón, (2014).

The study from Torres Calderón from 2014 combined two geophysical methods to generate a conceptual model of the geothermal system. Magnetotelluric (MT) data was collected using 70 stations and arranged in a grid of 150 m x 150 m x 15-100 m, while the Transient Electromagnetic Method (TEM) was applied using 36 stations. Figure 5 presents the final inversion model combining the two methods showing resistivity cross-sections. These are interpreted to show a

high-temperature system with a low-resistivity cap above a high-resistivity core. Isotherms were sketched using the TEM resistivities and geochemical data.

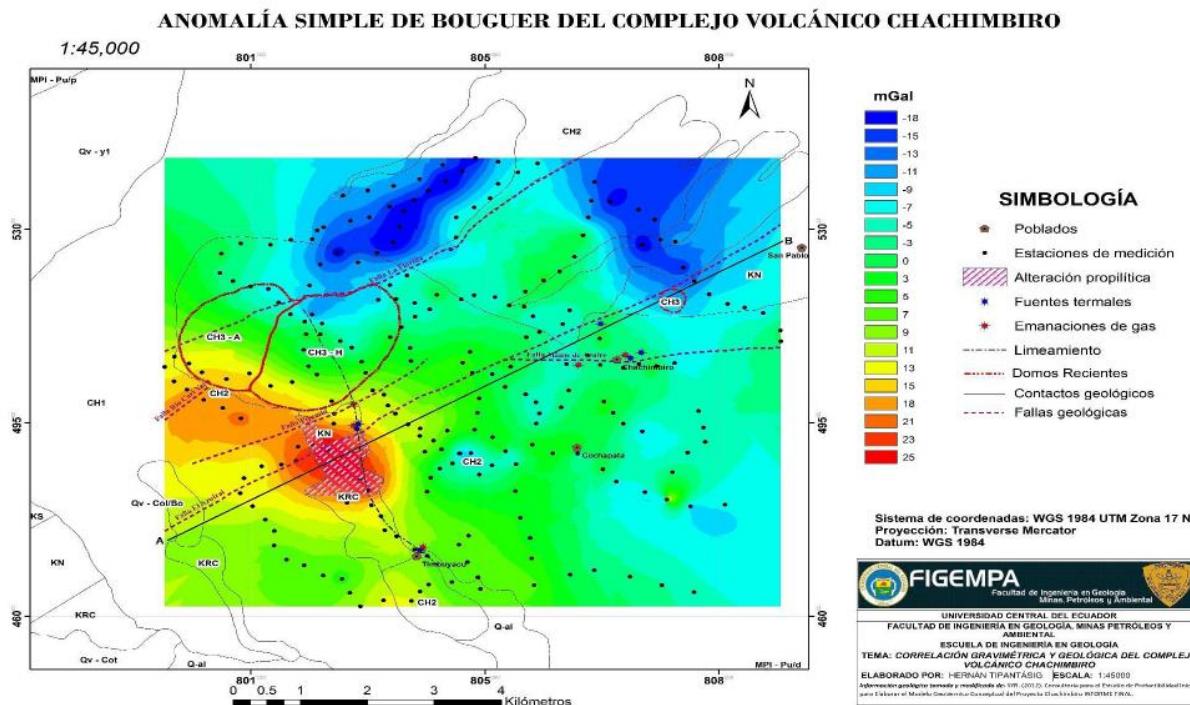


Figure 6: Correlation of the gravimetric results with structural and geological information from the Chachimbiro volcanic Complex, image from Córdova Tipantásig, (2017).

Córdova Tipantásig (2017) used the gravimetric method by measuring relative gravity from 234 stations georeferenced with a GPS. The parameters taken into account for the study were location, topography, accessibility and a distance between stations of ~250 m. The northern area shows an abrupt change between low and intermediate gravimetric values, while, the middle part shows changes between high and intermediate values. This changes in the gravimetric values may indicate a density change of related rocks and the presence of a lineament NE-SW direction and other with NW-SE direction.

Figure 6 shows a simple Bouguer anomaly map, suggesting that high anomaly areas could be related to high density from basement rocks with hydrothermal alteration. Also, the author suggests that an intrusive body is related to previous volcanism of the volcanic complex. High densities can be related to propylitic alteration generated by the fluids that filled rock cavities, reducing its permeability and increasing cohesion. On the other hand, intermediate gravimetric values could be related to lower densities from a mixture of low cohesive materials from debris avalanche of

the Chachimbiro 1 and 2. Finally, the author takes into consideration geological data from the surface. From this, they interpret low gravimetric values to be related to low density materials, likely sedimentary rocks and volcanic deposits, suggesting the basement is located at great depth.

Pilicita (2014) used Magnetotelluric (MT) data collected in 2011 by WesternGeco Integrated EM CoE. In total 70 stations were located on an irregular grid with a spacing ~ 0.35 km between them. The study was carried with eleven profiles in total. Several cross-sections were done with the data collected. The author identifies three anomalies separated from each other by a fault, two of them cover large areas. A shallow anomaly may correspond to the cap rock and covers 11.07 km^2 with an average thickness of ~ 388 m.

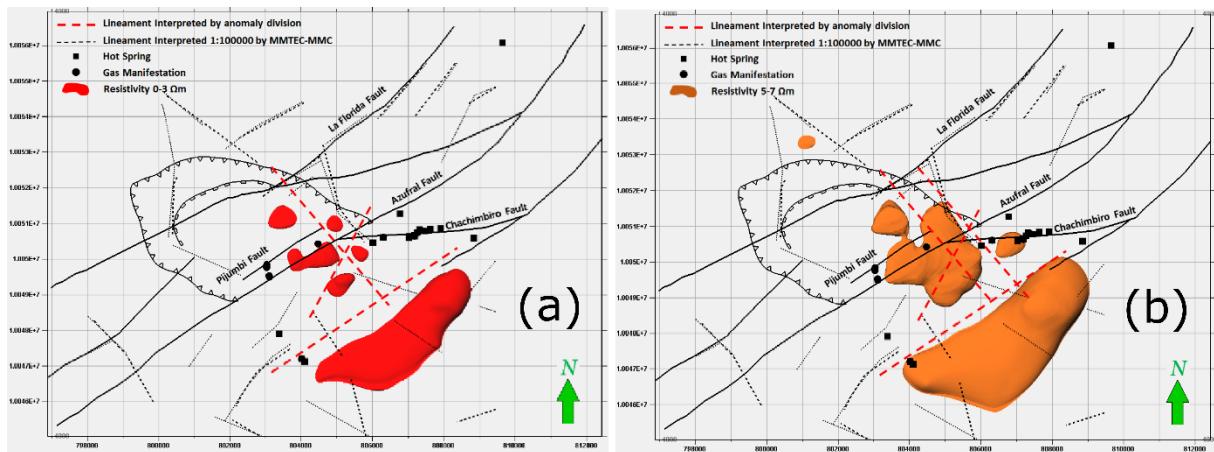


Figure 7: Low resistivity anomalies, faults and fractures in the Chachimbiro geothermal area. a) Resistivity anomalies between 0 to 3 Ωm . B) Resistivity anomalies between 5 to 7 Ωm . Image from Pilicita (2016).

Figure 7 (a) show the moderately low resistivity anomalies appear to be offset or separated by the Azufral and Pijumbi faults; this anomaly may be the cap rock. While (b) present the Azufral, Chachimbiro, Pijumbi faults and a lineament passing through the anomalies. A lineament in NW-SE direction is shown close to the borders of 5 to 7 Ωm resistivity anomaly. Lineaments increase permeability of the system and outer boundaries. Another lineament is subparallel to the Azufral fault that divides two anomalies. Resistivity distribution shows several layers; a shallow part with a high resistivity of $>160 \Omega\text{m}$, corresponding to Quaternary volcanic rocks and tills; a lower resistivity zone $<10 \Omega\text{m}$ which are possibly related to the cap rock, formed by smectite, and has been described in other technical reports. Below this low resistivity anomaly, resistivity values increase from 10 to 160 Ωm , forming a medium resistivity zone that includes a high resistivity core of 60 – 160 Ωm .

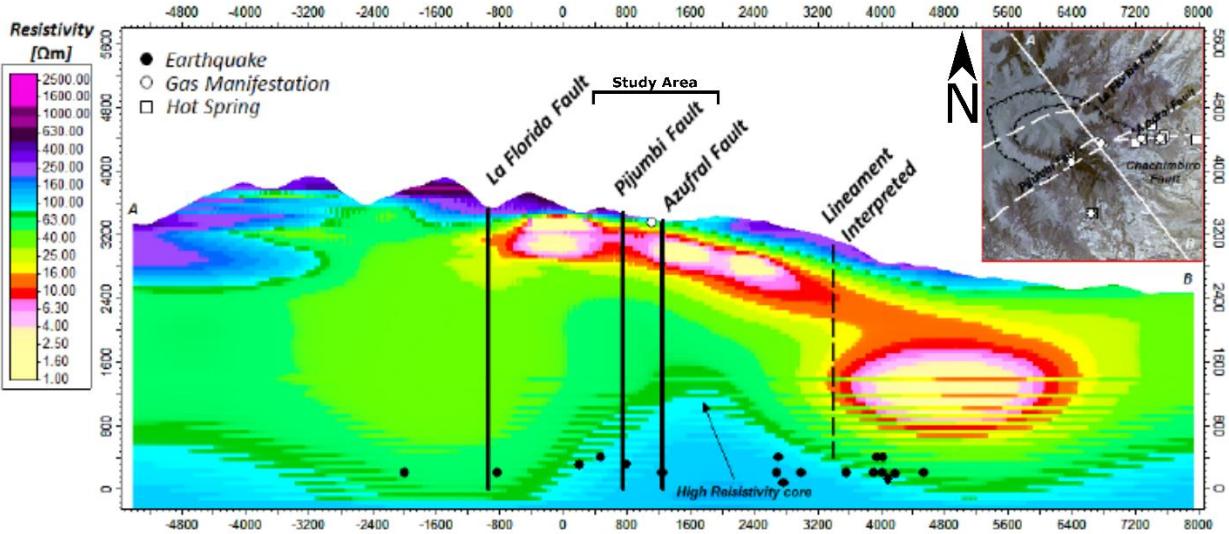


Figure 8: Resistivity cross-section showing the high resistivity core and the relation between the interpreted lineament, earthquakes, fault, gas manifestation and resistivity anomalies. Figure from Pilicita (2016).

Figure 8 shows the cross-section from the MT data collected, with a projection of the faults, reaching the high resistivity core. Hot springs and gas emissions are located along the faults, indicating that the core's fluids have connection with the surface through the faults.

Previous studies helped to create several models of the Chachimbiro's geothermal system, which allowed them to estimate the approximate depths of the reservoir and the cap rock. Córdova Tipantásig (2017) determined an approximate depth of the cap rock around ~600 m, while the reservoir top is approximately at ~2000 to ~2100 m. The Torres Calderón (2014) model shows the estimated depth of the cap rock is around ~300 to ~600 m, and reservoir top at ~1500 m depth. The Pilicita (2016) model shows a similar estimated depth of cap rock around ~300 to ~600 m with a thickness ~388 m, and a reservoir top at ~1500 m depth. As for the heat source, Bernard et al., (2014) showed magmatic reservoirs at ~14.4 and 8 km depth.

5. Methods

5.1. Magnetic theory

Magnetic fluxes exist in a magnet, going from the positive pole (north pole) to the negative pole (south pole), both poles having the same intensities but different sign. The direction and intensity of the magnetic field vary from point to point at the Earth's surface because the Earth behaves on average as a giant magnet. The force F between two magnetic poles of magnitude m_1 and m_2 separated by a distance r is given by:

$$F = \frac{\mu_0 m_1 m_2}{4\pi\mu_R r^2} \quad (1)$$

Where μ_0 is the magnetic permeability of free space which has a value of $4\pi \times 10^{-7}$ [WbA⁻¹m⁻¹] and μ_R denotes relative permeability, that is the ratio of the permeability of a specific medium to the permeability of free space.

$$\mu_R = \frac{\mu}{\mu_0} \quad (2)$$

In terms of magnetic susceptibility

$$\mu_R = 1 + k \quad (3)$$

Where k is the magnetic susceptibility, which is dimensionless

The force is attractive when the poles have different signs and repulsive if the poles have the same sign. μ_R is dimensionless.

The magnetic field B is a consequence of the current flux lines between two poles per unit area.

The units of magnetic field are the Tesla ($T = \text{Wb}^1\text{m}^{-2}$). The magnetic field is given by:

$$B = \frac{\mu_0 m}{4\pi\mu_R r^2} \quad (4)$$

The magnetic field is represented by force lines and its direction and intensity correspond to the number of lines per unit area.

The magnetic moment is the strength and orientation of a magnet or any other object that has a magnetic field, given by:

$$m = pr \quad (5)$$

Where m is a vector with the direction of unit vector r that extends from a negative magnetic pole to a positive magnetic pole p

When a material is located within a magnetic field, it can acquire a magnetization in the same direction as the applied magnetic field, which it will lose if the object is removed from the applied field, which is lost if it is removed. This phenomenon is called induced magnetization or magnetic polarization and result from the dipole alignment to the field direction within a material. The induced magnetic intensity J_i of a material is defined as the sum of all dipole moments m per volume unit V , given by:

$$J_i = \frac{\sum m}{V} \quad (6)$$

The induced magnetization intensity is proportional to the magnetization force H of the inducing field:

$$J_i = kH \quad (7)$$

Where H is the surrounding magnetic field.

The total magnetic field including the magnetization effects results in:

$$B = \mu_0(H + J) \quad (8)$$

Combining equation (7) and (8) results

$$B = \mu_0(1 + k)H \quad (9)$$

Replacing it with (2), we get

$$B = \mu_0\mu_R H \quad (10)$$

5.2. Remanent magnetization in magnetic survey

Rocks and minerals can show a remanent or permanent magnetization when after removing them from an applied field. Remanent magnetism in igneous rocks results from magnetization of rocks

during their formation, with cooling of these rocks below the Curie temperature leading to acquisition of the magnetic field of that time. The Curie temperature is the value at which the minerals lose their magnetic properties. Magnetic susceptibility is the measure of the magnetic field in response to an external field (Harrison et al., 2015). All materials have magnetic susceptibility, which could be diamagnetic or paramagnetic, ferromagnetic, antiferromagnetic. Diamagnetic materials (e.g. marble, salt, quartz) have all their atomic orbitals full of electrons. When an external field is induced, electrons align in the opposite direction of the applied field. This results in a low and negative susceptibility. Paramagnetic materials (e.g. platinum, aluminum, pyroxenes, biotite) have incomplete atomic orbitals, resulting in an alignment in the same direction of the external field applied. The field produced by these materials is relatively low. Ferromagnetic materials (e.g. cobalt, nickel, magnetite) have positive susceptibility. They show a strong spontaneous magnetization that could exist even in absence of an external field, which can decrease with rising temperature or be completely lost if the Curie temperature is reached. This value varies depending on the mineral composition. Antiferromagnetic materials (e.g. hematite) have divided magnetic domains aligned in opposite directions, with the same quantity of moments in the same directions which results in a magnetic field equal to 0. Ferrimagnetic materials (e.g. garnet, cubic ferrites of iron oxides with other elements such as cobalt, nickel, manganese, zinc) are aligned in opposite directions to the external field, but the magnetic moment is different from 0. This happens because some subdomains have different magnetic components than others. This effect allows a high magnetic susceptibility and a spontaneous magnetization. Figure 9 shows the difference between ferrimagnetism, ferromagnetism and antiferromagnetism.

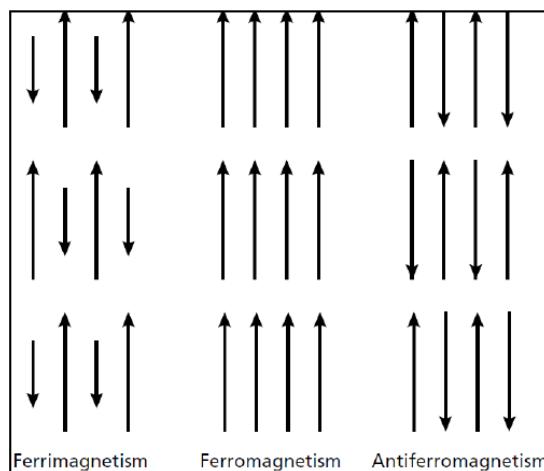


Figure 9: Magnetic moment direction scheme, figure from López Males et al. (2013).

5.3. Magnetic surveying

Magnetic surveys have different applications ranging from micro-scale, archeological studies (detection of metallic objects) to regional-scale studies. It can be used on the ground, air, or sea (I et al., 2013). The objective of a magnetic survey in a geothermal area is to acquire information about the relationship between geothermal activity, tectonics, and stratigraphy. This will be done with detection and interpretation of the magnetic anomalies of underground rocks and magnetic properties of the ground (Sircar et al., 2015). Magnetic measurements in geothermal exploration are commonly used to detect the location of intrusions, dykes, faults and possibly estimate their depth. It is also possible to locate areas with lower magnetization due geothermal activity (Bjornsson, 1980).

All rocks are magnetic, but, different types of rocks have different amount of magnetic minerals that will create anomalies. What one detects in a magnetometry survey are differences in magnetization of the subsurface rocks, which is a vector sum of the total induced and/or permanent magnetizations within the rock forming minerals. For example, if a ferrous material with an existing permanent magnetization is buried in the ground, it will also acquire an induced magnetization in the direction of the current Earth's magnetic field at that location and most likely in a different direction than the permanent magnetization resulting in a superimposed total magnetization. This will create a magnetic anomaly that is slightly different than what is expected at that location, if there were no magnetic material or if there were only magnetic material with no permanent magnetization. The detectability of such an anomaly depends on the amount of magnetic material, its size, the magnetization contrast between the object and the surroundings and its distance from the sensor. These anomalies can be produced by lithology changes, size of magnetized bodies, faulting and topography (Rivas, 2009). Figure 10 shows the result of the sum of vectors of remanent magnetization (M_r) and induced magnetization (M_i), giving as result the total magnetization (M_t).

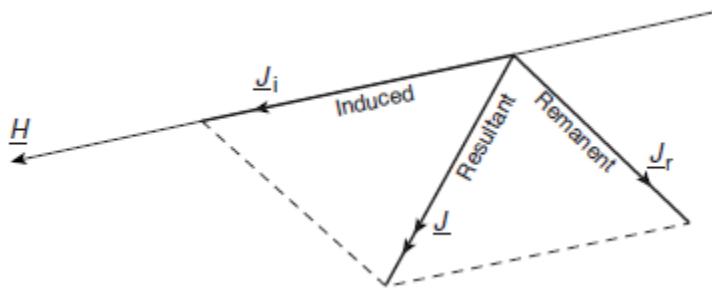


Figure 10: Addition of induced and remnant magnetization.

Several types of magnetometers exist for geophysical exploration, but the most known and the one used in this study is the proton free-precession magnetometer. This kind of magnetometer comprises a recipient filled with hydrogen-rich fluid like kerosene or gasoline that is surrounded by a coil made of copper wire (Figure 11a). The nuclei of hydrogen act like dipoles that are aligned with the Earth's magnetic field (Figure 11b). The equipment form a circuit with a current flowing through the coil and generating a magnetic field 50 – 100 times bigger than the surrounding magnetic field but with a different direction, forcing the protons to realign (Figure 11c) (I et al., 2013). After this current is turned off, the protons will align back parallel to the surrounding Earth's magnetic field (Figure 11d). However, this aligning back is not direct but happens through precession. The frequency of this precession and the time it requires is related to the strength of the magnetic field at that location.

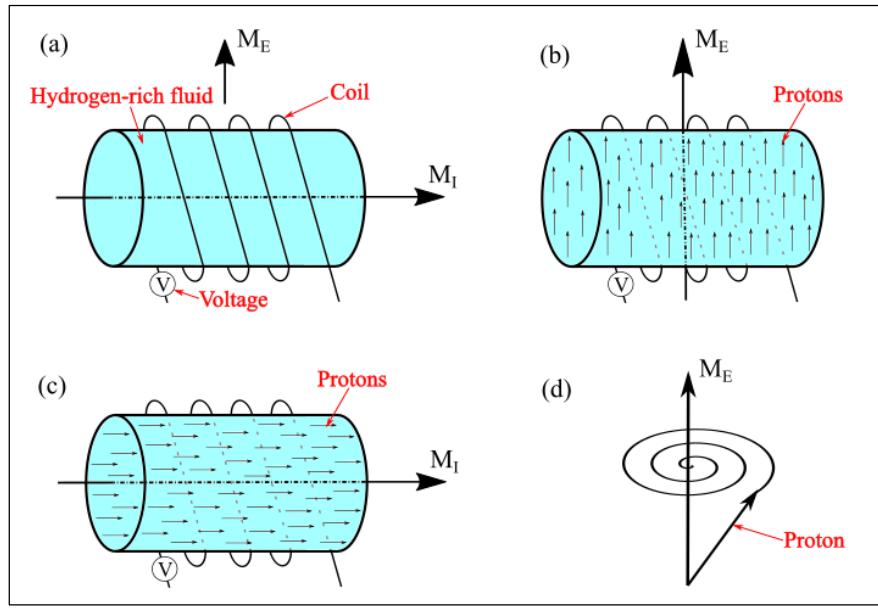


Figure 11: Principle of a proton-precession magnetometer: (a) Sensor with Earth's magnetic field M_E , and magnetic field of instrument coil M_I . (b) Alignment of protons in Earth's magnetic field. (c) Alignment of protons due to applied magnetic field. (d) Precession of protons around Earth's magnetic field after coil current is interrupted. Image from (Gallegos Aguilar, 2020).

5.4. Data collection

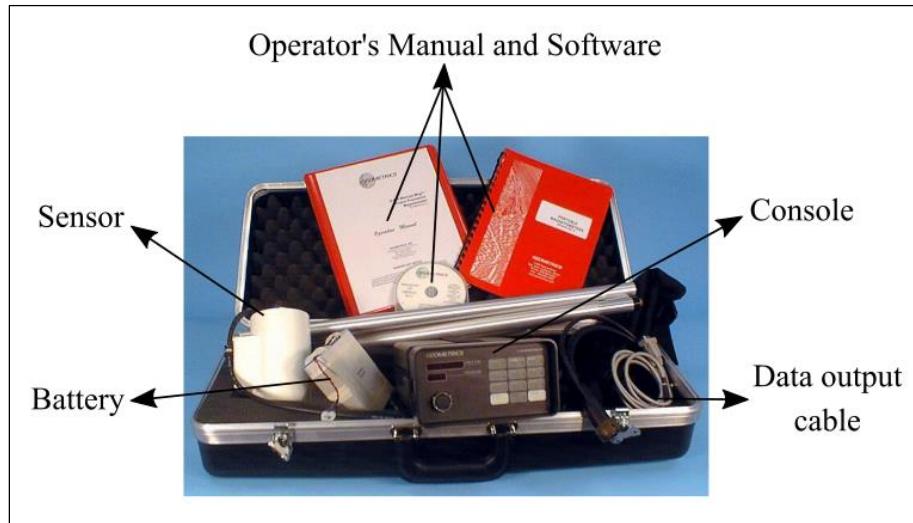


Figure 12: G-856AX proton-precession magnetometer used to take measurements. Image modified from Antennas Manual from Geophysical Survey System, Inc.

The equipment used for this research was a G-856AX proton-precession magnetometer that was provided by the Instituto Nacional de Patrimonio Cultural (INPC) (see Figure 12). This model is a single sensor magnetometer where the user needs to keep the body of the coil perpendicular to the surrounding Earth's magnetic field by approximately pointing the arrow of the magnetometer to the North and always handling the magnetometer in the same way to avoid errors in the measurements.

Magnetic values change during the day due to the interaction between the magnetosphere and the sun location, so corrections need to be applied to the collected data to account for these variations. To realize the corresponding diurnal correction, it is necessary to take "base station" readings every day before starting measuring and throughout the day at certain intervals depending on the size of the anomalies. In our case we were looking for large anomalies so a base station reading every 1 to 2 hours was sufficient to remove the daily variation. For optimal sensitivity the magnetometer was tuned every day to an average value of the magnetic field in the study zone, which in our case was 28500 nT. Then, data correction helped to reduce erratic values of the magnetic field (Telford et al., 1990).

The selected survey area is show in Figure 13. Field work was performed over a duration of one month from November 11th to December 3rd of 2020. The irregularity of the terrain and the presence of a dense forest prevented a regular grid to be applied for the survey. Instead, measurements were taken every ~27 to ~35 m, covering the areas that were the most accessible. Then, two more survey lines (Figure 13) were done across the faults, with a spacing between points of ~5 m. and one additional profile was extracted from the 30 m survey, for a total of three magnetic profiles.

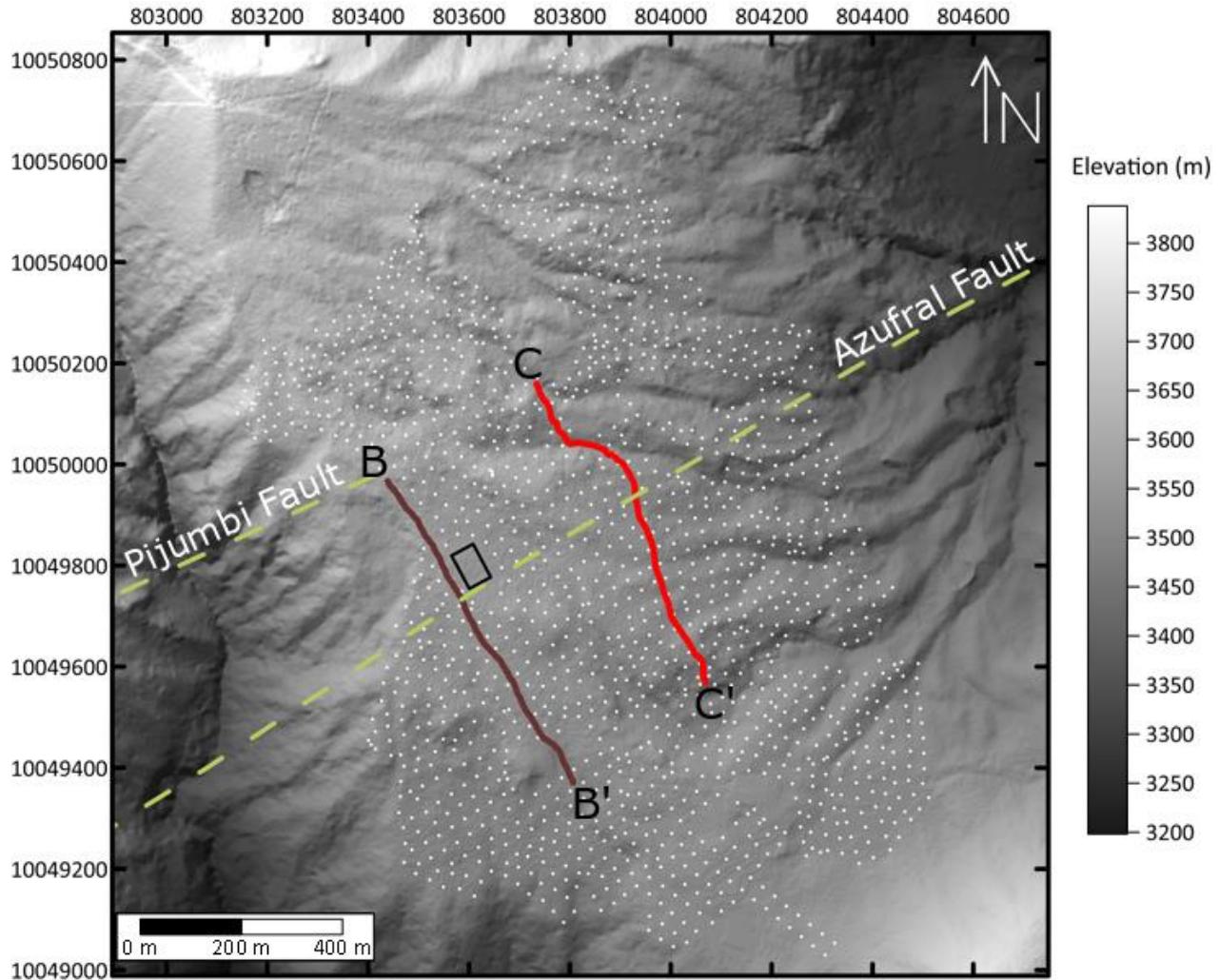


Figure 13: Location map showing the points acquired during the survey. B-B' in brown corresponds to the 5 m step size survey line 1 and C-C' in red corresponds to the 5 m step size survey line 2. Black rectangle is the area of the first drilling, also the two faults crossing the area are shown.

The measured data were noted manually and the location of each point recorded with a Garmin GPS. After fieldwork, the latitudes and longitudes were converted to Universal Transverse Mercator (UTM) coordinates, the base station correction was applied to every data point and the collected data were organized to the required format to be plotted in MagMap software for further processing. After despiking the data in Magmap, the files were exported to be plotted on Surfer GoldenSoftware as a magnetic anomaly map, after that the resultant map was set at the top of the topography map of the area to be compared then.

In addition, some models of a strike-slip fault were done using GravMagSuite app in MATLAB and a model with the clay cap using Mag2dc. This models would help to understand better the resultant magnetic profiles.

6. Results

Data was uploaded to Magmap2000 Software to have an initial filtering. The program shows the survey points' distribution with coordinates, and a profile of the magnetic data, and the user has the option to draw the data in 2D and/or 3D interpolating data in between the points in order to fill the grid. Figure 14 shows the raw data plotted as a color relief map in surfer. It is impossible to interpret, since the outlier values distort the color scale and will not show any details. These outliers are likely either typing errors, instrument errors or interference with unwanted magnetic objects while surveying. It was necessary to use the filter tool called range despike of Magmap2000 to remove these erroneous data that draws very high or very low anomalies that are out of range from the other values.

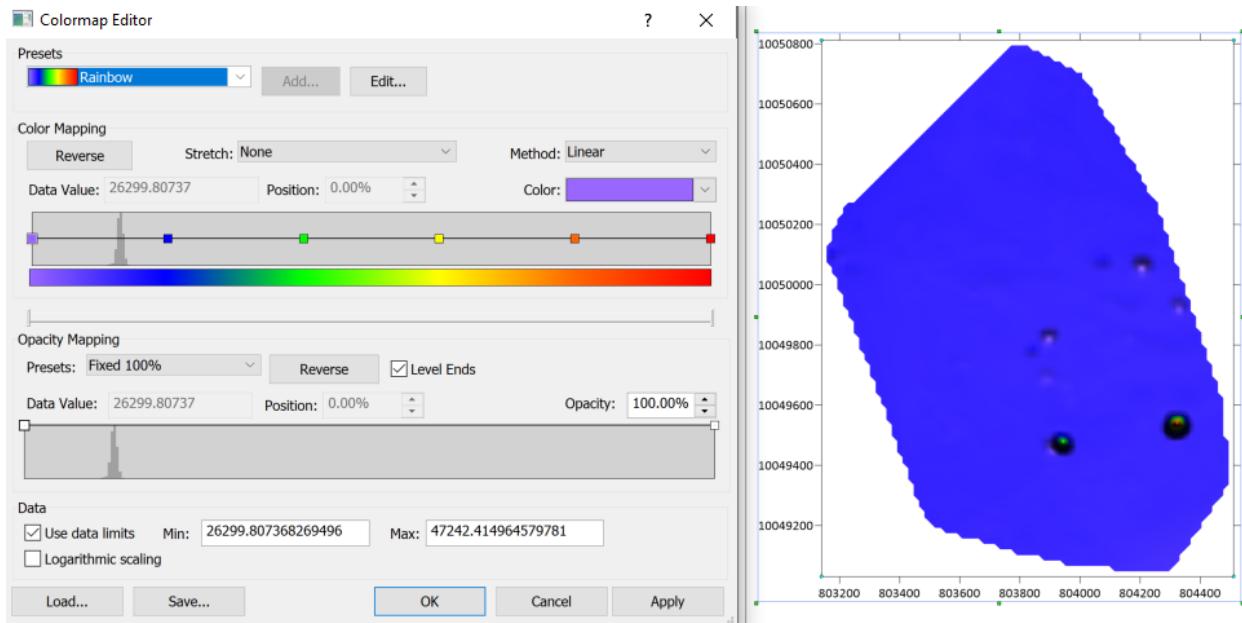


Figure 14: Raw data plotted without filtering in Surfer (i.e. despiking).

After data despiking, the data file was exported and plotted in Surfer GoldenSoftware, where, with the obtained results it was possible to create a magnetic anomaly map (Figure 15) of the study area. Also, some parts of the maps were drawn with help of data interpolation. The map shows different distribution of magnetic values in the area. The north part of the map is mainly characterized by high magnetic values and several positive magnetic anomalies, with the highest positive magnetic anomaly at the NW part of the magnetic map. The right southern part of the map

is characterized by low magnetic values with several negative magnetic anomalies. Finally, the left lower part is characterized by intermediate magnetic values, with no representative anomalies.

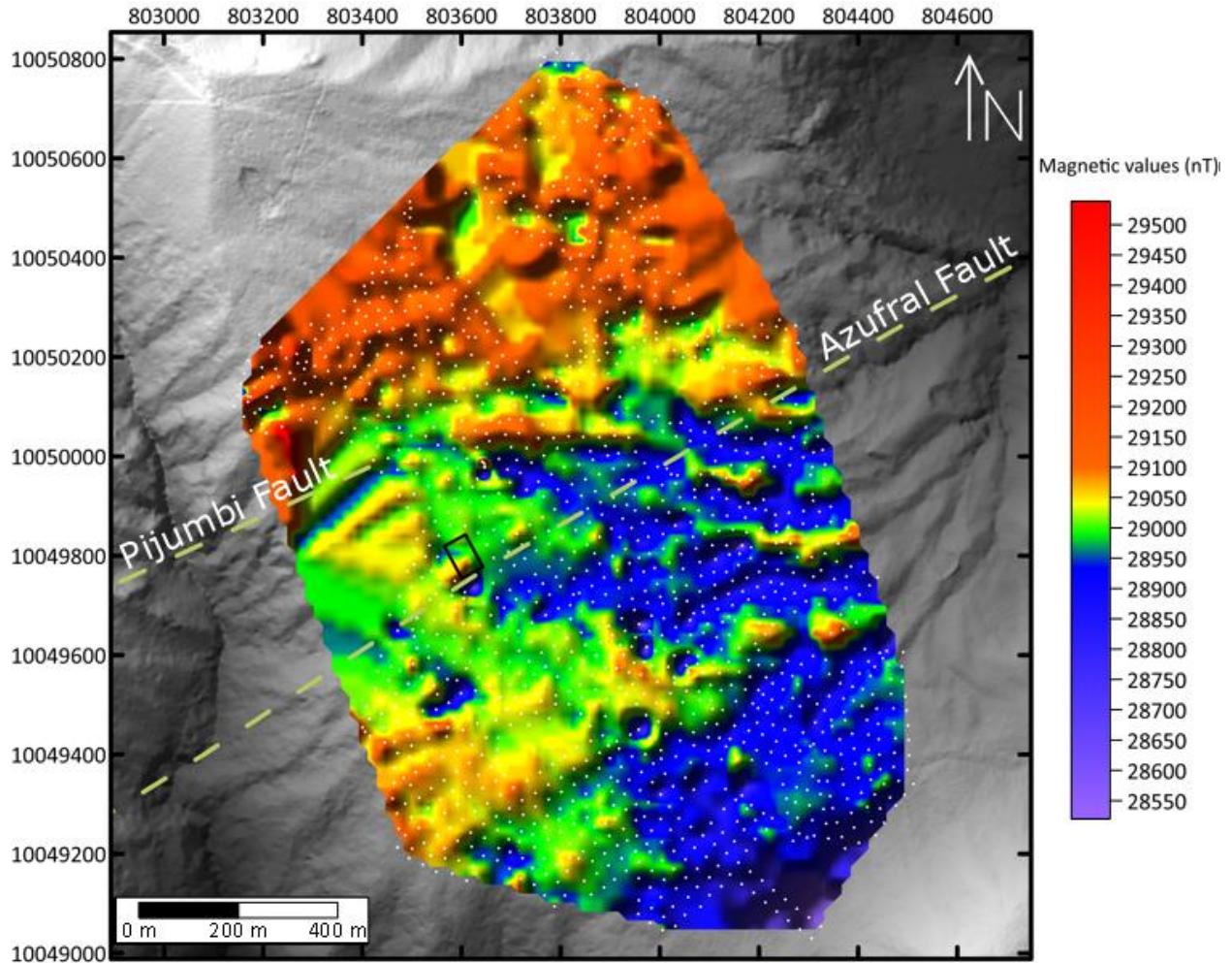


Figure 15: Magnetic anomaly map with survey points. Black rectangle is the area of the first drilling, also the two faults crossing the area are shown.

The resultant magnetic anomaly map was also plotted in 3D format (Figure 16), which helps to recognize the areas with positive or negative anomalies. Then, it is possible to compare the survey data with the study area's elevation map, discussed in the next section.

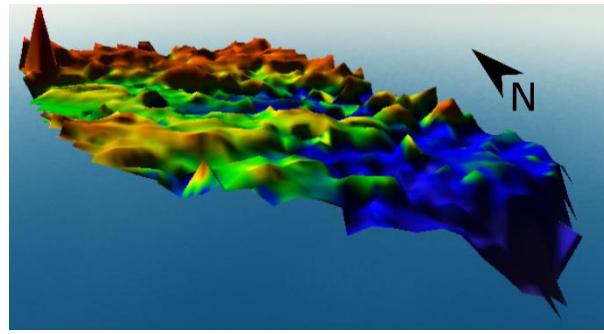


Figure 16: 3D magnetic anomaly map made in Surfer, view from South to North.

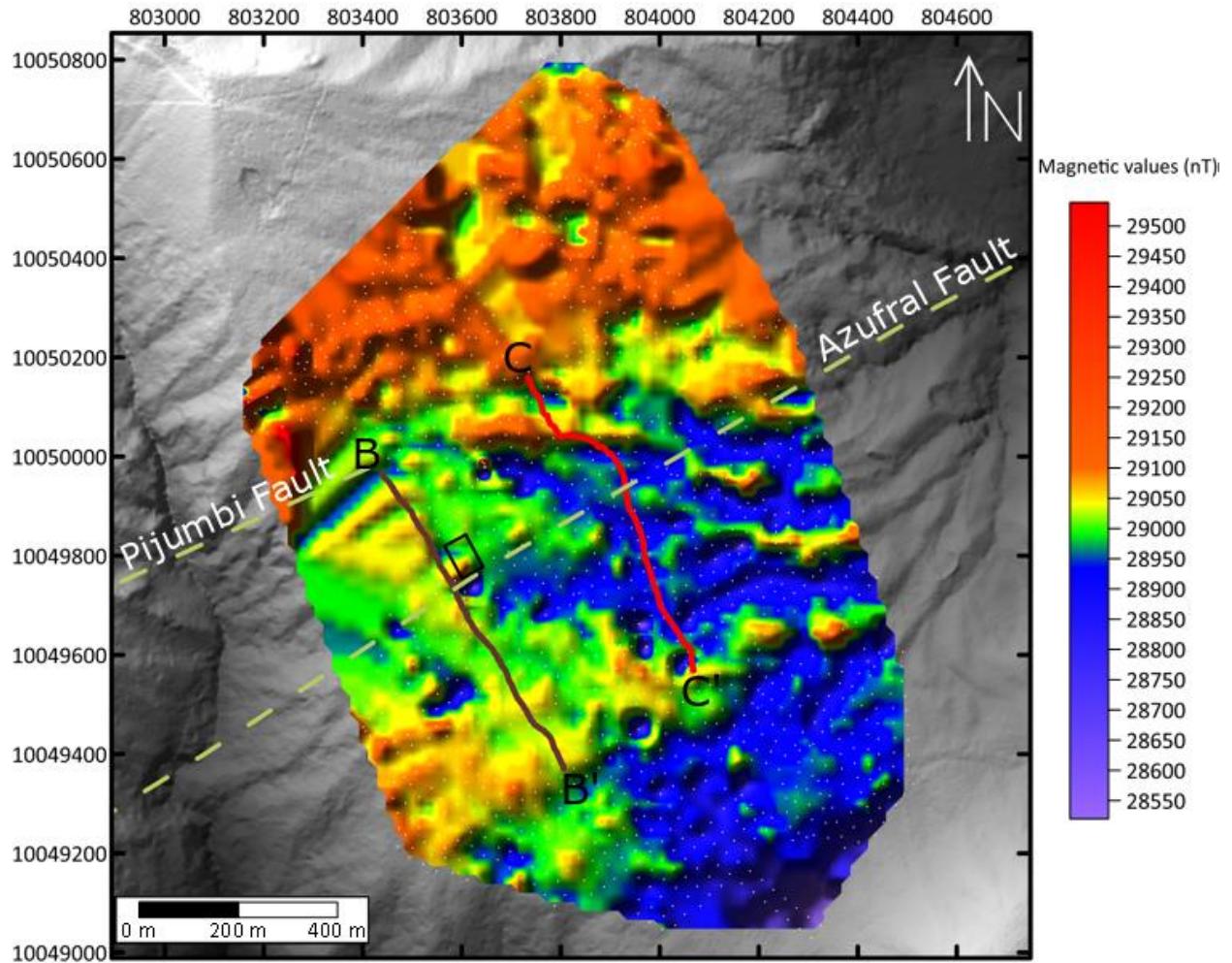


Figure 17: Magnetic map with profile lines. B-B' in brown and C-C' in red. Black rectangle is the area of the first drilling, also the two faults crossing the area are shown.

In addition to the 2D map, two profiles were extracted from the elevation and magnetic maps (Figure 17); the resulting elevation and magnetic profiles are shown in Figure 18.

In the case of B-B', the elevation profile (Figure 18a) is continuously decreasing from NW to SE, while the 5 m magnetic profile (Figure 18b) initial values increase from values under 28900 nT, then rises and fluctuating between 29100 and 29000 nT at the interval from 0 to 100 m. Then at the interval from 100 to 200 m, values abruptly decrease under 28900 nT, only to rise again between 29100 and 29000 nT. Values fluctuates again between 29100 and ~28950 nT; next to that, values abruptly decrease close to ~28950 nT and increases between 29200 and 29100 nT at the interval 200 to 300 m. After that values keeps fluctuating between 29100 and 29000 nT; close to the end of the profile values increase again close to 29100 nT at the interval 600 to 700 m. The 30 m survey profile (Figure 18c), similar to the last profile values start a values close to 28950 nT from the interval 0 to 100 m. Then values rise close to 29050 nT with a value decreasing of ~50 nT. Next to that values increase again ~50 nT in next interval from 100 to 200 m, followed an abruptly decrease to low magnetic values at interval from 200 to 300 m. After that it keeps fluctuating between 29050 and 29000 nT. Magnetic data from this study is shown on the Appendix chapter (Table A1).

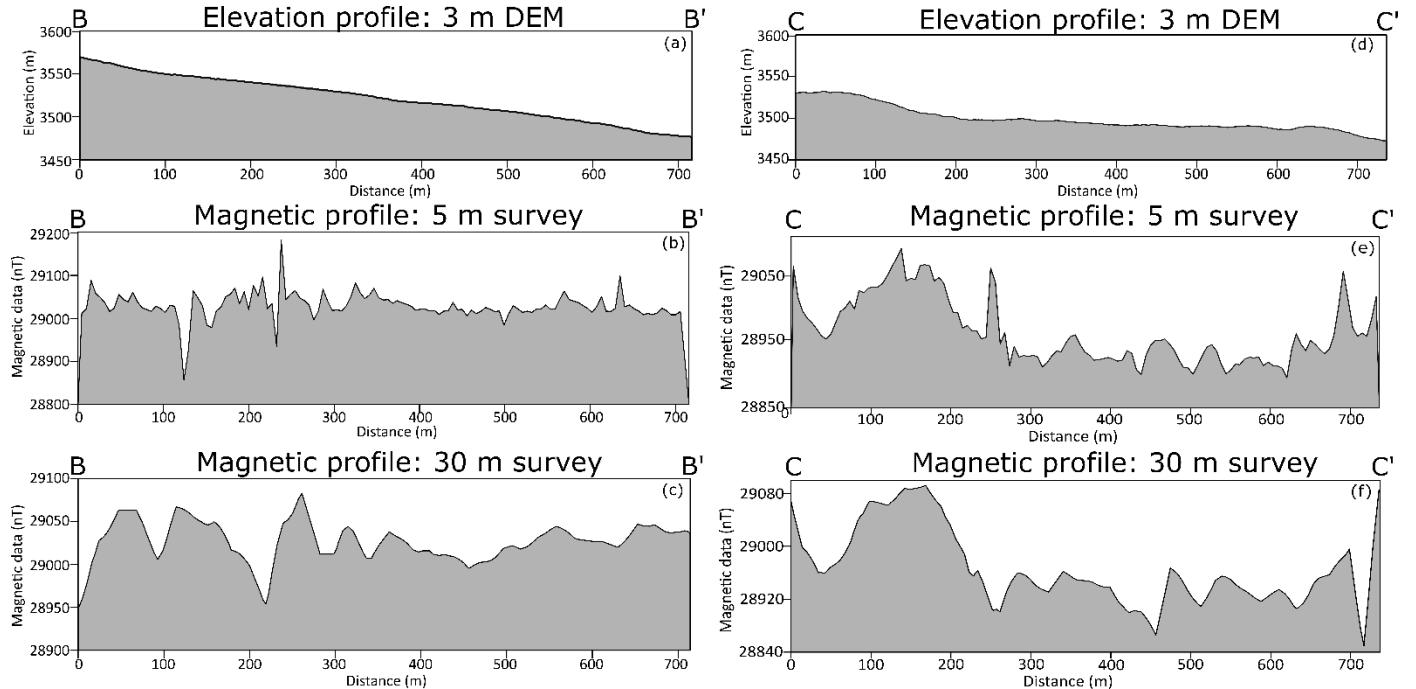


Figure 18: Elevation and magnetic profiles. B-B' correspond to the fault survey line 1; C-C' correspond to the fault survey line 2. Profile lines are shown in Figure 17.

For line C-C', it is possible to observe a little change in the elevation profile, starting with high values and then decrease continuously (Figure 18d). For the magnetic profile for 5 m survey

(Figure 18e), start with high magnetic values, fluctuating a little until it drops a little under 29000 nT at the interval from 0 to 100 m, then, values increase reaching values above 29050 nT. Then values decrease at the start of interval from 200 to 300 m with an abrupt increase in between the interval. Followed values fluctuate in between 29000 and 28900 nT. At the final part of the profile at interval 600 to 700 m, value abruptly rise again. For the magnetic profile for 30 m survey (Figure 18f), at the interval from 0 to 100 m values start at ~29060 nT; next to that values decrease close to ~28960 nT, then values rise again close to 29080 nT. After that, the profile keeps decreasing until it reaches values less than 28920 nT. Values then rise again ~40 nT at the interval from 200 to 300 m. Values continue fluctuating a little with a subsequent decrease again reaching values in between 28920 and 28840 nT, later values rise again in between 29000 and 28920 nT at the interval from 400 to 500 m. Finally, values fluctuate close to 28920 nT until it abruptly decreases close to 28840 nT to rise again at ~29080 nT after the interval from 600 to 700 m.

7. Discussion

7.1. The effects of faulting and hydrothermal alteration on magnetization

According Yang et al. (2020), temperature, stress, and fluids in a fault system may affect iron-bearing minerals because its sensibility to them. Their magnetic properties give an interpretation of the physical and chemical process that affects rocks on faults. Faults go through cycles of creep and seismic slip. A seismic cycle is divided into three periods: 1) coseismic period (Figure 19a), which is the time during an earthquake during seconds to minutes; 2) postseismic period (Figure 19b) which time could be days, months and even years after an earthquake; finally, 3) the interseismic period (Figure 19c) which happens between large earthquakes, taking tens to thousands of years.

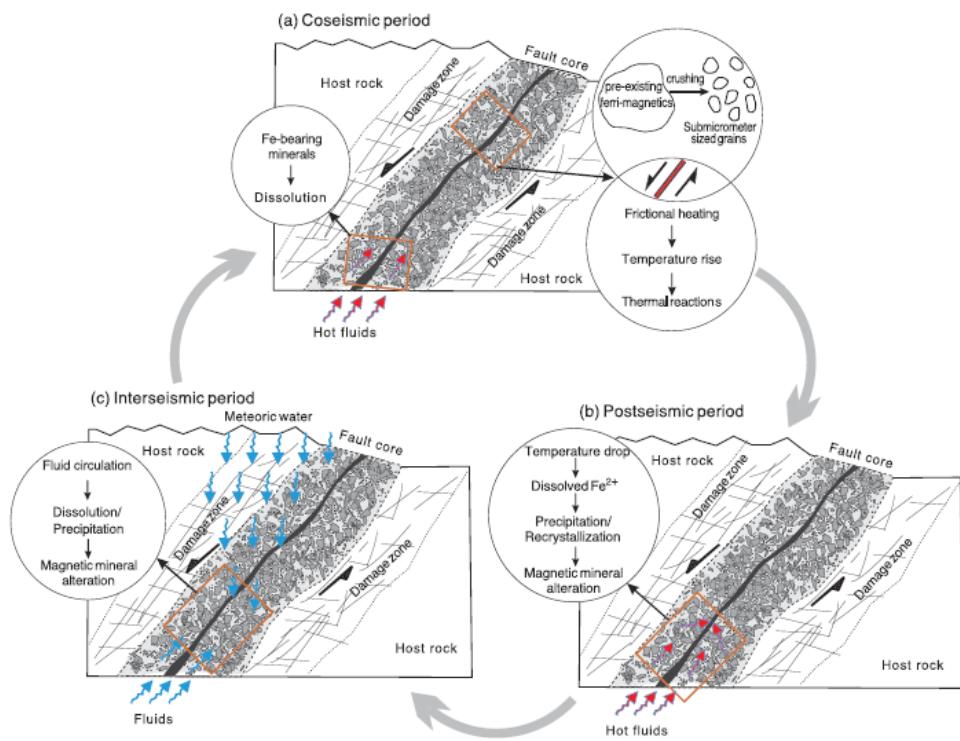


Figure 19: Conceptual model showing the faulting-related physical and chemical processes and the causes of potential magnetic changes in the fault zone during different stages of the earthquake cycle. (a) Coseismic period, (b) Postseismic period and (c) Interseismic period. Image from Yang et al., (2020).

Yang et al. (2020) describe the processes that occur in active fault zones during the seismic cycle, which influence the magnetic properties of the rocks at fault zones. Thermochemical reactions (Figure 19a) take up the largest part of the total energy from an earthquake (~80% to ~90%).

Frictional heat will raise the temperature in the slip zone of rupture. With typical seismic rates (1 m/s) and total slip distance (tens of centimeters to meters) at a fault plane, temperature rises <100 °C at shallow depths close to the surface but it can increase to >1100 °C at higher depths (>5 km). This increased temperature may allow dehydration of some mineral phases and formation of breakdown products; the increased temperature at the slip zone sometimes is enough to melt the host rock minerals. Temperatures inside a geothermal system do not reach Curie temperatures, where minerals get entirely demagnetized. However, hydrothermal alteration may affect mineral magnetization by reducing it or replacing into less magnetic minerals.

Another process that affects the magnetization is the fluid movement in faults: fault zones are formed by dense networks of fractures and secondary faults which act as fluid conduits in the crust. Fluids from several sources, such as meteoric waters, mineral dehydration, trapped formation brines, and volatiles from deep underlying layers infiltrate and percolate along fault zones. Pore fluid pressurization during frictional heating may allow fault weakening. These fluid-related dissolution-precipitation processes are common during all periods of the seismic cycle. Resultant reactions play an important role in physical, chemical, and mechanical evolution of fault rocks.

Magnetic anomalies are useful to delineate high-temperature hydrothermal/geothermal systems; it is possible because several geothermal processes may alter the rock magnetization, going through a demagnetization of minerals or altering them to a less magnetic mineral (Caratori Tontini et al., 2016). Temperatures inside a geothermal system do not reach Curie temperature, where minerals get entirely demagnetized. However, hydrothermal alteration may affect minerals magnetization reducing it or replacing into less magnetic minerals. Magnetic anomalies may be caused by intrusions, flows, or iron-rich sediments. The magnetic method is useful to locate intrusive bodies, tracing individual buried dykes and faults or estimating their deep (Bjornsson, 1980).

7.2. Forward modelling

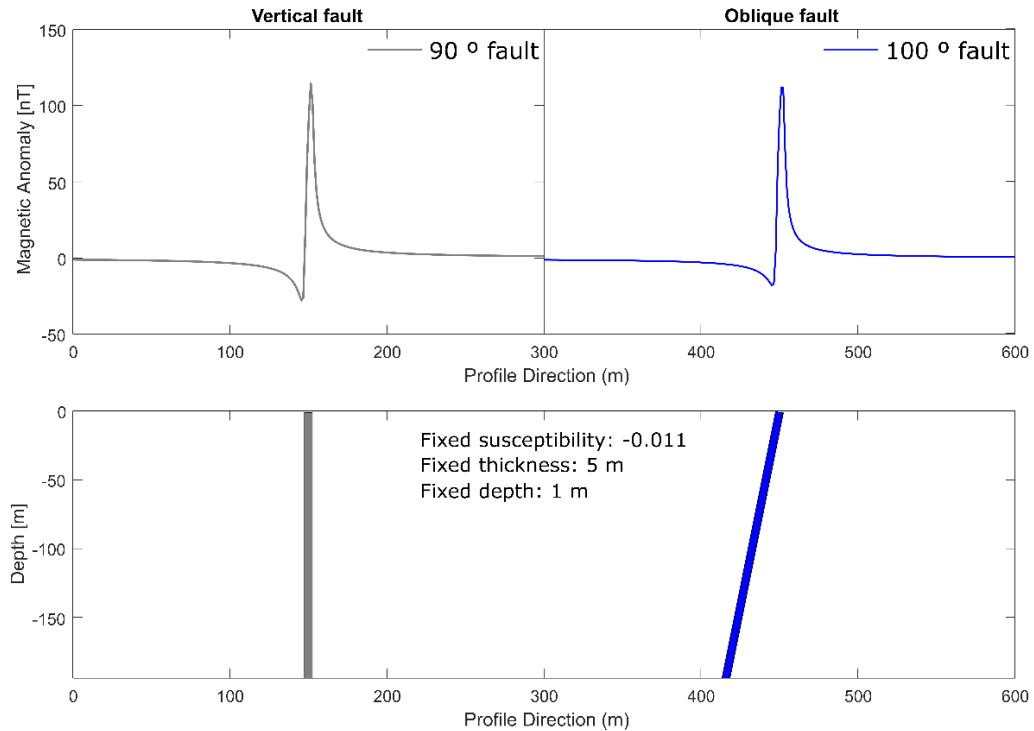


Figure 20: Strike-slip fault model comparison with 90 ° and 100 ° inclination. Model created with dike-like forward modelling tool from GravMagSuite app.

Models were created at MATLAB with GravMagSuite app, using the modelling feature for dike-like forward modeling tool. Values for inclination of 21.5441 °, a declination of -4.2683 ° and a total field of 28859.5 nT correspond to the study area. It was used fixed susceptibility values for andesite, described in (Li and Fu, 2019). Also, other fixed values such as the fault inclination, depth and thickness were used to see it similarity with major anomalies in the profiles. The resultant models help to understand what kind of effect would the fault present in a magnetic survey. The first model (Figure 20) represent the behavior of an strike slip fault, one with a 90 ° and other with 100 ° inclination. This models show an abruptly variation of magnetic values from low to high values around ~120 nT. Both of them have similar behavior but the 100 ° has an slightly variation of ~10 nT in comparison with the 90 ° fault.

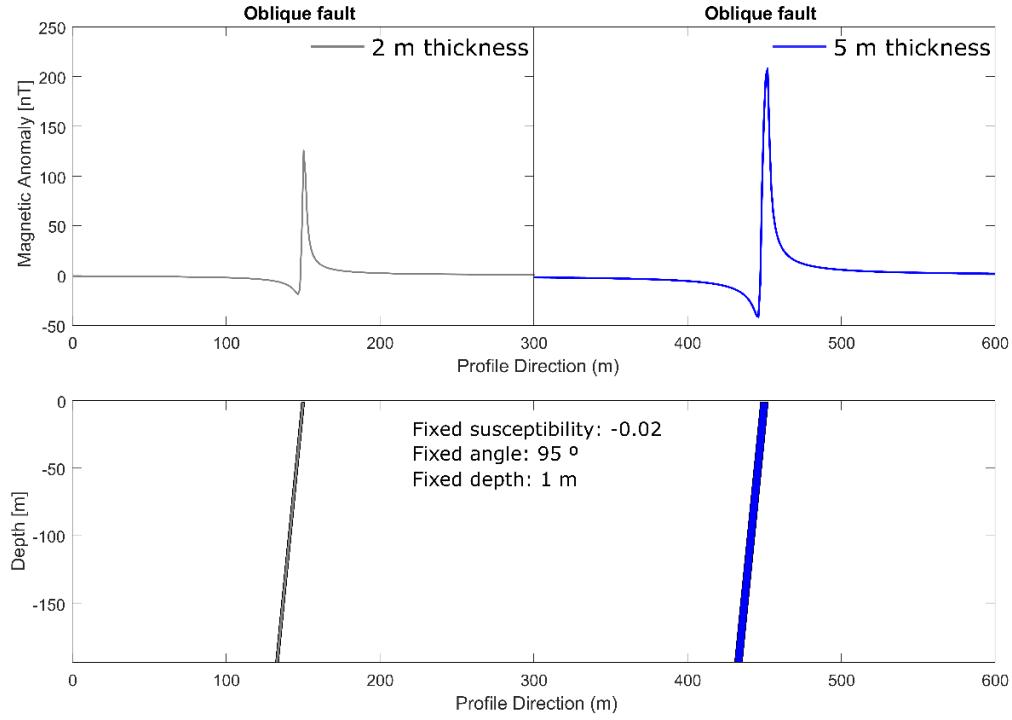


Figure 21: Strike-slip fault model comparison with 2 and 5 m thickness. Model created with dike-like forward modelling tool from GravMagSuite app.

While, the second model (Figure 21) was created with the same total field, inclination and declination configuration, but using a fixed susceptibility, depth and inclination values. In this case, the model compares the effect that the thickness of the fault would have. The first faults show a thickness of 2 m while the second show a thickness of 5 m. Both of the models show to have a similar form for the anomaly but the 2 m thickness show an abruptly change from low to high magnetic values around ~ 120 nT, while the 5 m thickness show a higher variation in data around ~ 220 nT.

Finally, the third model was created with the same total field, inclination and declination configuration, but using a fixed susceptibility, thickness and inclination values. In this case, the model compares the effect that the depth of a fault would have. The first model uses a 2 m depth for the fault showing a very abrupt change in the magnetic values around ~ 200 nT, while the second model uses a 10 m depth, showing a much less effect with a value change around ~ 50 nT. These models would help to interpret anomalies showed in the magnetic profiles.

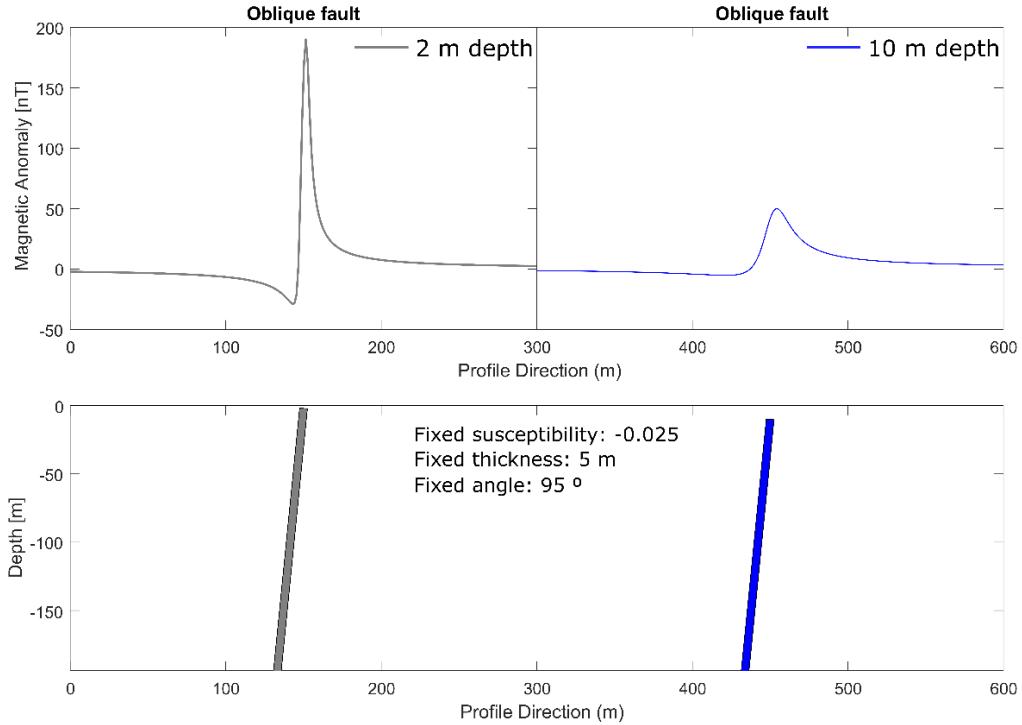


Figure 22: Strike-slip fault model comparison with 2 and 10 m depth. Model created with dike-like forward modelling tool from GravMagSuite app.

The next model (Figure 23) again uses the same configuration of the area for the inclination, declination and magnetic field. This one represents what would be the effect of a clay cap, but using a fixed susceptibility value of -0.05 for andesite to see the effect that a similar body would have. Magnetic values tend to decrease when they are close to the effect of the clay cap as it is expected and increase while they get far from it. Andesite value used is described in (Glen et al., 2007).

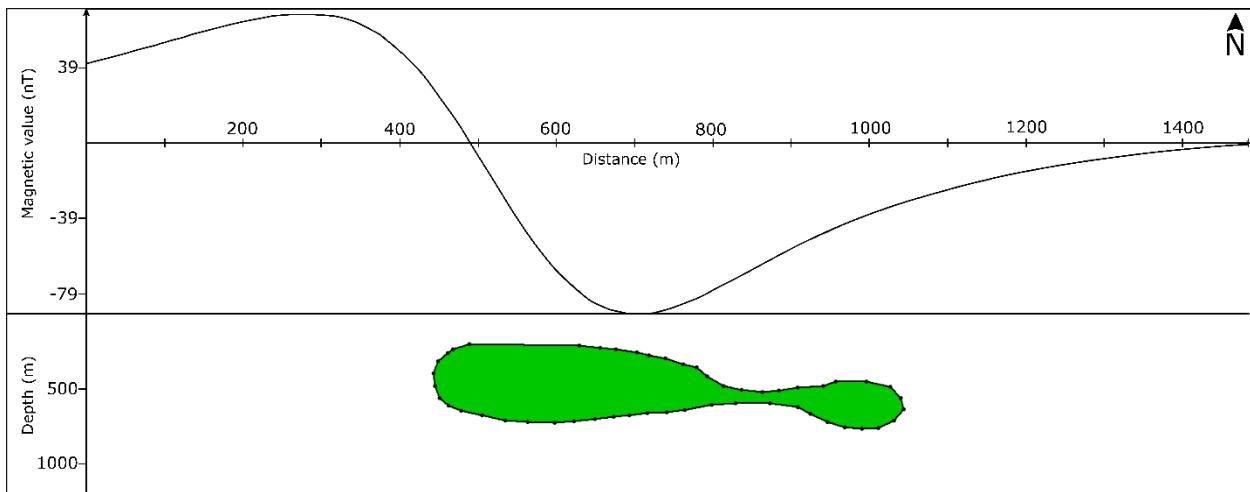


Figure 23: Clay cap model, created on Mag2dc software.

7.3. Interpretation of magnetic anomalies

A geological map of the area is shown in Figure 24A with the major geological units in the zone. These units can be compared with the magnetic anomalies in the study area (Figure 24B). The northern part of the area is characterized by the presence of positive magnetic values with a major positive magnetic anomaly at the NW part of the study area. The whole Chachimbiro's area is comprised of volcanic materials associated with high magnetic values due to its high magnetite content (Kearey et al., 2002); but, this part of the zone may be associated more with materials from the domes Albuji and Huga, composed of dacitic rocks. Likewise, the northwestern area does not show much hydrothermal alteration that may affect rock magnetization and this area comprises the highest elevations registered in the survey.

The western part of the map is associated with the Tumbatu unit, which is composed of pyroclastic flows, ignimbrites, tephras and landslide deposits. This area shows intermediate values with no significant magnetic anomalies; this could be because the area is now covered by pyroclastic deposits of later eruptions of Chachimbiro. These materials have less magnetic susceptibility in comparison to the primary products of a volcanic eruption. Also, the area may be affected by hydrothermal alterations close to the fault zone. Hydrothermal alterations could be caused by carbonization, related to CO₂ emissions; propylitic intermediate temperatures, which may reduce the magnetic properties of the andesites and basaltic andesites, and argillic alteration, associated with hot springs and fumaroles from the current geothermal activity.

Finally, the central eastern and SE parts are characterized by low values and negative magnetic anomalies. These values may be related to hydrothermal alterations, carbonization, related to CO₂ emissions, and steam heated advanced argillic alteration. These alterations are present in a large area along the Chachimbiro and Azufral faults. Also, the values follow the river valleys, where influence by topography on ground magnetics is very apparent in stream gorges (Telford et al., 1990). Additionally, the lower values could be related with thicker sedimentary deposits filling the area.

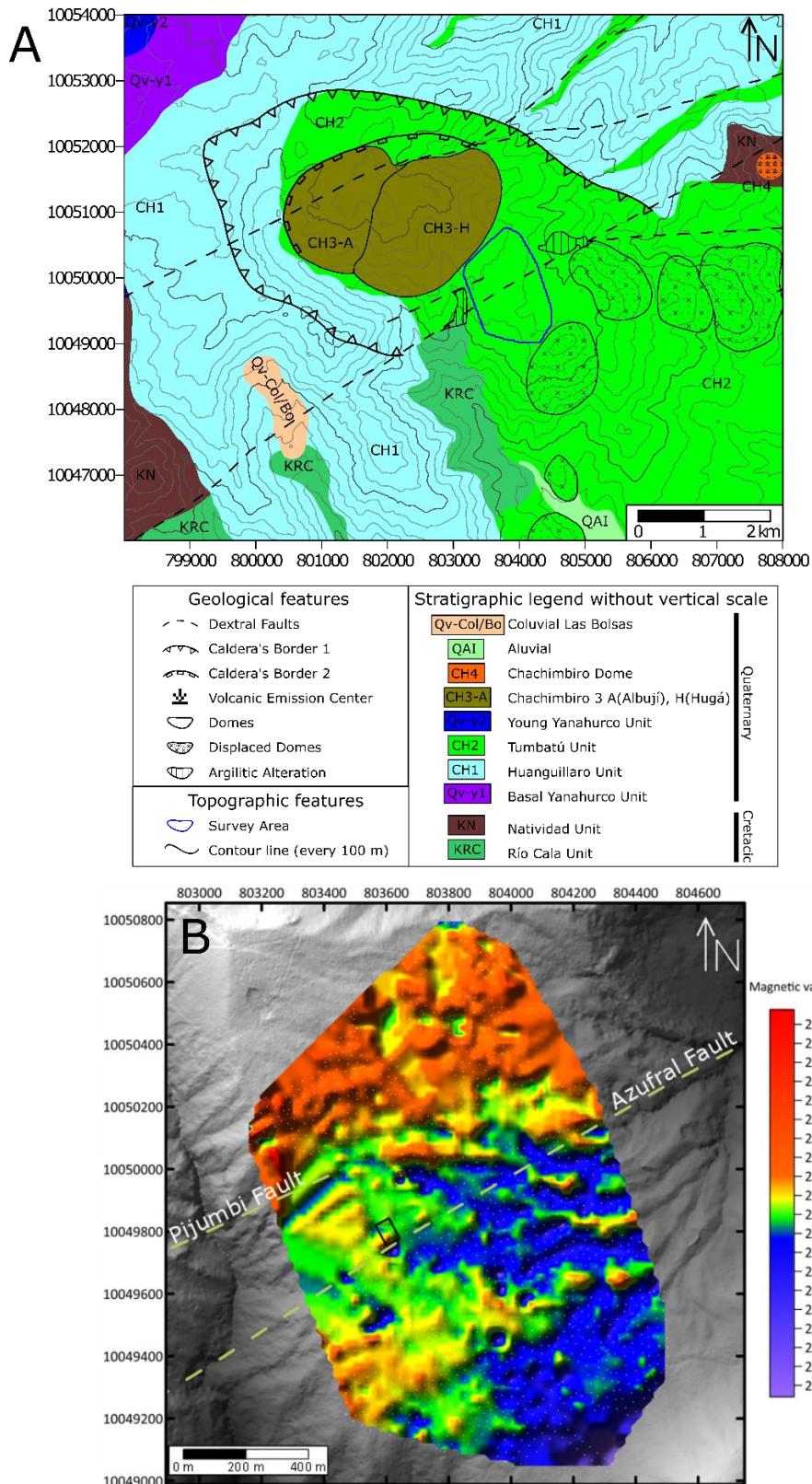


Figure 24: Comparison between the geological map and the magnetic anomaly map. (A) Geological map of Chachimbiro volcanic complex, modified after (SYR, 2012). (B) Resulting magnetic anomaly map.

7.3.1. Topographic effect on magnetic surveys

Surfer GoldenSoftware gives the option to create profiles from both elevation and magnetic values from the map, allowing us to see how much the area's topography may affect the magnetic values. The influence of topography in magnetic surveys can be significant, but it is not completely predictable; after diurnal corrections, the variations in magnetic values may only depend on the magnetic properties of topographic features (Kearey et al., 2002). Similar to a dike or fault effect, topography effect will have an associated anomaly but much smaller in comparison to them (Ugalde et al., 2013). In this case, several profiles were created, from survey lines using 5 m spacing survey and one extracted from the 30 m survey.

Figure 24A show the magnetic profiles with some areas circled in black, which correspond to the anomalies that could be related to a topographic effect in the survey, while Figure 24B show the areas related to that effect, circled in black. For B-B' profiles there is only one change that could be related to the topographic effect. In the case of the 5 m survey, the change circled in black seems to be more exaggerated while in the 30 m survey, it is not so appreciable. This difference in the values could be because the 30 m survey does not cover much of the part of the area where the B-B' passes.

For C-C' profiles there are several areas circled in black. For the first circled area, bot profiles show similar changes from low to high magnetic values, corresponding to a change in topography in the area circled in Figure 24B. Values fluctuates more in the 5 m survey while in 30 m are smother, this could related to the distance used to measure the magnetic values. For the next area circled in black at interval from 400 to 500 m, value change from low to high magnetic values of ~60 nT is observed in both profiles, this caused probably by the influence of the quebrada. Finally the last area circled in black at the interval after 700 m, shows and abrupt change to high values and then a decrease to lower values, in the case of the 30 m survey the decreasing is more exaggerated than the 5 m survey, this change could be some values in between the 30 m survey could be interpolated with other distant values with a high difference in topography in comparison to the continuous measuring of 5 m survey.

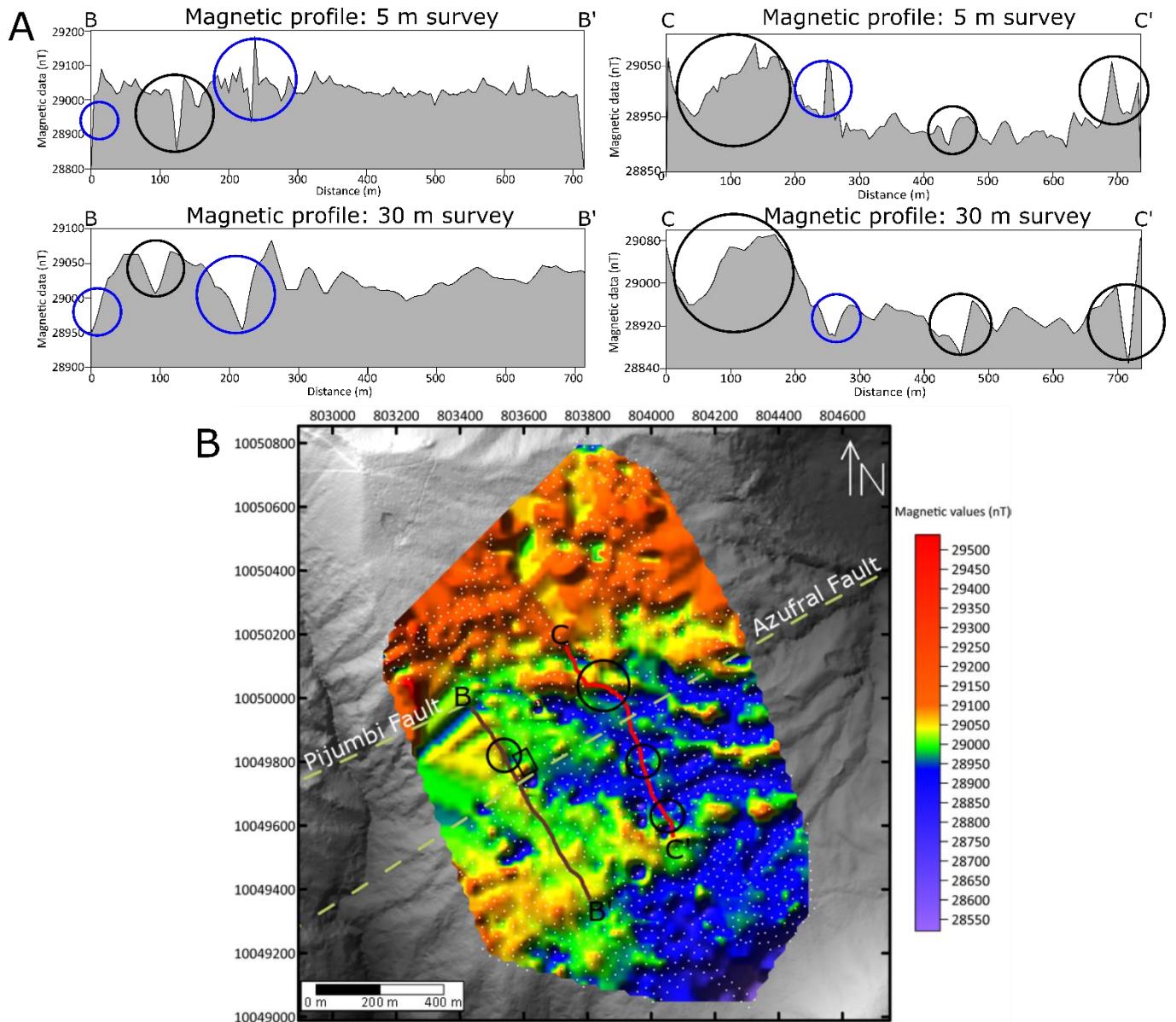


Figure 25: Magnetic profile comparison with the resulting magnetic map. (A) Corresponds to the magnetic profiles. Anomalies in blue circles show the fault effect in magnetic data while black circles the topographic effect. (B) Profile lines B-B' in brown from 30 m survey and C-C' with black circles. Black circles remark the location where topography may affect the magnetic values.

7.4. Faults detection in magnetic survey

Figure 25A shows two abrupt changes circled in blue that may represent the effect of fault in the magnetic survey. First magnetic anomalies circled in blue for B-B' profiles show a similar behavior with the 2 m depth fault model, the 2 m thickness fault model and fits with any of the models showing the inclination effect in a fault. With a difference between low and high magnetic

values around ~150 nT. These anomalies and the models named before may represent the interaction of the Pijumbi fault with the magnetic values at the start of both B-B' profiles. While, the next blue circles at both B-B' profiles from 200 to 300 m interval show a similar behavior with 2 m depth fault model, 5 m thickness fault model and any models showing the inclination effect in a fault; with a difference between low and high magnetic values around ~200 nT. These anomalies, due it similarities with the models named before may represent the interaction of the Azufral fault.

For C-C' line it is only possible to see one abruptly change that could be related to a fault. In the case of the 5 m survey the blue circle shows a differentiable abrupt change from low to high magnetic values, while in the 30 m survey the area with the blue circle is not so exaggerated. In the case of the 5 m survey we can relate the anomaly can be compared with any of the models showing the effect of fault inclination, the 2 m depth fault model and the 2 m thickness fault model, while in the 30 m survey the same, except that for this one shows more similarities with the 10 m depth fault model. This difference between the two profiles could be related to the resolution of the survey, the topography related to the quebrada and the presence of the clay cap that is concentrated in the right and right-lower part of the area.

The study showed to be effective in detecting magnetic anomalies in the area, most of the negative anomalies being located in the SE part of the map. We interpret that the low values correspond to the clay cap rock, previously described by Pilicita (2016) that shows that the resistivity anomalies could be related to the presence of the cap rock at the southeast part of the study area. Guillén, (2020) described an estimated depth for a smectite layer around ~200 m and an illite layer around ~250 m; while, Torres Calderón (2014) and Pilicita (2016) studies estimate the clay cap rock at an approximate depth of ~300 to ~600 m.

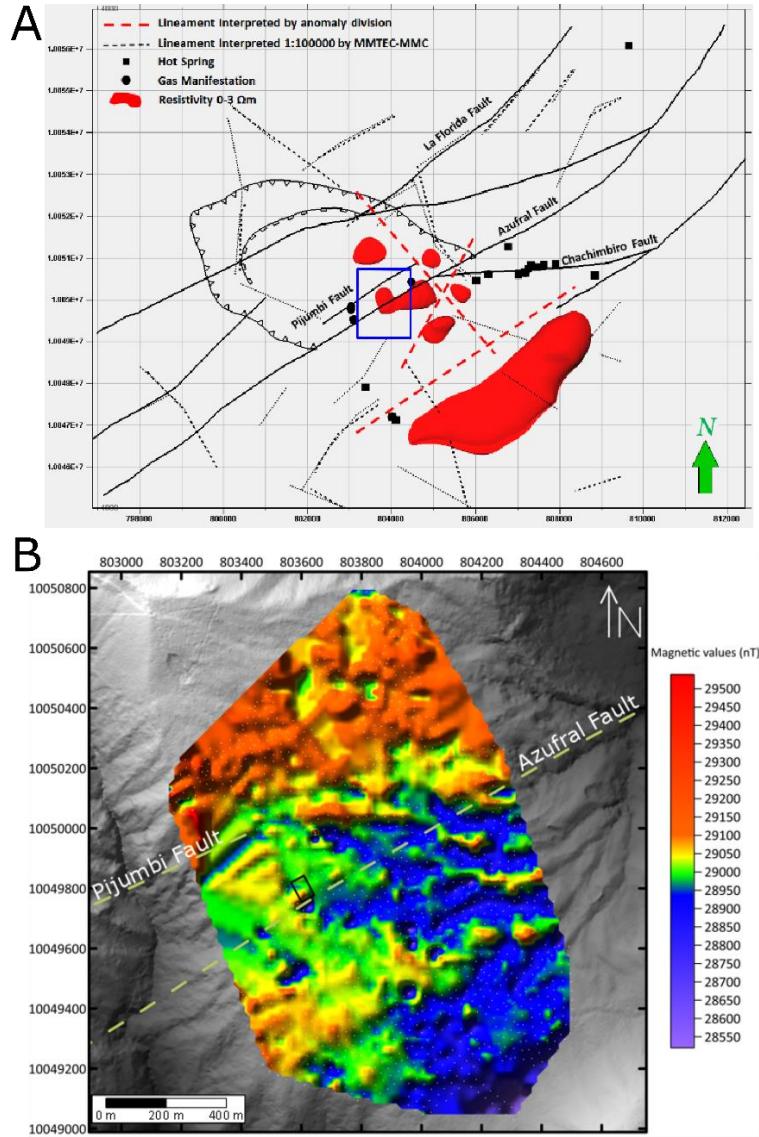


Figure 26: Comparison of the map with previous studies. (A) Resistivity anomaly map from Pilicita (2016), blue rectangle represent the study area. (B) Magnetic anomaly map from this survey.

The magnetic behavior of the geological units is also reflected in the magnetic map. Areas closer to mapped geothermal alterations show a reduced magnetization as expected. The distribution of magnetic anomalies also fit with the resistivity model presented by Pilicita (2016) in Figure 26A, which shows the a resistivity anomaly which correspond to the location of the clay cap at the right most part of the survey area; while in the magnetic map Figure 26B, the right and right-lower part show the major concentration of low magnetic anomalies as is expected for the effect of the clay cap in magnetic values. Finally, faults location and effect seems to slightly offset compared to the area where it was previously mapped.

8. Conclusions

This study used magnetometry applied to geothermal exploration in the Chachimbiro area, using two different approaches: a large scale grid survey with ~30 m spacing, and two high-resolution lines with ~5 m spacing. The survey of the area resulted in a magnetic anomaly map where it is possible to see the distribution of the positive and negative anomalies along the area. Also, elevation and magnetic profiles were compared from ~30 m survey and ~5 m survey lines, showing that topography affects magnetic values; positive anomalies are related to high topographic values and negative anomalies are related to low topographies. Fault effects on the magnetic survey are observable in the magnetic fluctuation values from the profiles, showing the fault effect in magnetization, while the other abrupt changes could be more related with topographic effects. Also, it is possible to see the demagnetization effect of hydrothermal alterations in the area after comparing with the mapped geological units. Comparison with previous model show that the area with major quantity of low magnetic anomalies could be related to the location of the clay cap of the geothermal system. The study shows that it is possible to use the magnetic method in geothermal areas to see the effect of faulting, topography and geothermal alterations on rock magnetization. This study could be used in future studies to contribute with the estimated location of the clay cap in the geothermal area.

To improve the results of this survey, a good option for future work would be to do another magnetic survey taking less resolution, which may help to see if it fits with data from the present work and the fault effect across the area. Also, it would be good to use different geophysical methods, such as resistivity survey along the fault area and see if it shows similar results where the faults are located in this and previous studies.

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10. Appendix

Table A1: Measured magnetic data with their coordinates.

Station ID	Latitude	Longitude	UTM_X	UTM_Y	Magnetic_data (nT)
1	0.452391	-78.275829	803237.715	50059.76	29539
2	0.453715	-78.275558	803267.849	50206.279	29470
3	0.452838	-78.276443	803169.298	50109.197	29452
4	0.454074	-78.276233	803192.64	50245.977	29354
5	0.451724	-78.272167	803645.683	49986.105	29353
6	0.453548	-78.275541	803269.749	50187.801	29299
7	0.453918	-78.276052	803212.809	50228.722	29288
8	0.453672	-78.276014	803217.053	50201.502	29286
9	0.453355	-78.276512	803161.59	50166.403	29276
10	0.452929	-78.276106	803206.835	50119.281	29255
11	0.45337	-78.275681	803254.161	50168.098	29255
12	0.45385	-78.276198	803196.548	50221.191	29254
13	0.455834	-78.273483	803498.911	50440.847	29254
14	0.457576	-78.269275	803967.603	50633.789	29252
15	0.455123	-78.269916	803896.299	50362.321	29250
16	0.453831	-78.275852	803235.093	50219.103	29247
17	0.453636	-78.275751	803246.352	50197.529	29238
18	0.453177	-78.27541	803284.358	50146.753	29233
19	0.452936	-78.276389	803175.309	50120.044	29230
20	0.454949	-78.274336	803403.926	50342.88	29227
21	0.453443	-78.275422	803283.01	50176.187	29227
22	0.456945	-78.269973	803889.873	50563.935	29226
23	0.454238	-78.273551	803491.403	50264.237	29224
24	0.458046	-78.269702	803920.016	50685.779	29222
25	0.453252	-78.275605	803262.632	50155.044	29220
26	0.456767	-78.270911	803785.389	50544.198	29220
27	0.453979	-78.275224	803305.044	50235.507	29218
28	0.456482	-78.273511	803495.765	50512.551	29214
29	0.455723	-78.270964	803779.528	50428.671	29210
30	0.458141	-78.269863	803902.076	50696.285	29208
31	0.456736	-78.269547	803937.338	50540.826	29208
32	0.453445	-78.275942	803225.083	50176.386	29207
33	0.458292	-78.268856	804014.249	50713.037	29205
34	0.455649	-78.273901	803452.354	50420.358	29204
35	0.455346	-78.273469	803500.491	50386.847	29204
36	0.453118	-78.276006	803217.967	50140.199	29203
37	0.452642	-78.27632	803183.008	50087.514	29202
38	0.454119	-78.274066	803434.038	50251.047	29201
39	0.45417	-78.273819	803461.551	50256.701	29201
40	0.453663	-78.275041	803325.444	50200.547	29200
41	0.454131	-78.274327	803404.962	50252.364	29200
42	0.453938	-78.274245	803414.105	50231.011	29200
43	0.455824	-78.273622	803483.427	50439.735	29199
44	0.454491	-78.273842	803458.975	50292.221	29199
45	0.45618	-78.269116	803985.374	50479.319	29198
46	0.45347	-78.276191	803197.344	50179.142	29197
47	0.455698	-78.271127	803761.372	50425.897	29195
48	0.455478	-78.270197	803864.981	50401.592	29194

Station ID	Latitude	Longitude	UTM_X	UTM_Y	Magnetic_data (nT)
579	0.450032	-78.273231	803527.225	49798.83	29044
580	0.453419	-78.270416	803840.671	50173.741	29044
581	0.448808	-78.27255	803603.138	49663.415	29044
582	0.447173	-78.271502	803719.951	49482.536	29044
583	0.446315	-78.270319	803851.771	49387.641	29044
584	0.44577	-78.272735	803582.655	49327.234	29044
585	0.445483	-78.271241	803749.096	49295.537	29044
586	0.45356	-78.269312	803963.649	50189.39	29043
587	0.451343	-78.272093	803653.942	49943.948	29043
588	0.447256	-78.272487	803610.22	49491.679	29043
589	0.447768	-78.269932	803894.823	49548.441	29043
590	0.453779	-78.268925	804006.752	50213.64	29042
591	0.453846	-78.26748	804167.72	50221.115	29042
592	0.452371	-78.269862	803902.43	50057.796	29042
593	0.451468	-78.271952	803669.644	49957.786	29042
594	0.447707	-78.269484	803944.732	49541.71	29042
595	0.453988	-78.268831	804017.214	50236.771	29041
596	0.452715	-78.272429	803616.455	50095.755	29041
597	0.449181	-78.273406	803507.766	49704.654	29041
598	0.449762	-78.273556	803491.032	49768.94	29041
599	0.447283	-78.273775	803466.738	49494.614	29041
600	0.446738	-78.272131	803649.899	49434.374	29041
601	0.447906	-78.267909	804120.177	49563.796	29041
602	0.450262	-78.265757	804359.81	49824.593	29041
603	0.453659	-78.268349	804070.922	50200.386	29040
604	0.452015	-78.272297	803631.189	50018.301	29040
605	0.44588	-78.271006	803775.258	49339.477	29040
606	0.452791	-78.273404	803507.839	50104.124	29039
607	0.450384	-78.270076	803878.673	49837.913	29039
608	0.450896	-78.27179	803687.715	49894.497	29039
609	0.446165	-78.270545	803826.601	49371.034	29039
610	0.446677	-78.270237	803860.891	49427.702	29039
611	0.447529	-78.269588	803933.154	49522.009	29039
612	0.446378	-78.267603	804154.328	49394.725	29039
613	0.45178	-78.274172	803422.327	49992.218	29038
614	0.453596	-78.268248	804082.176	50193.419	29038
615	0.447256	-78.271771	803689.982	49491.709	29038
616	0.446346	-78.271941	803671.081	49391.005	29038
617	0.45263	-78.266416	804286.299	50086.601	29038
618	0.450255	-78.266134	804317.813	49823.803	29038
619	0.455188	-78.27264	803592.847	50369.399	29037
620	0.447392	-78.274038	803437.436	49506.664	29037
621	0.446765	-78.27424	803414.959	49437.275	29037
622	0.447358	-78.269693	803921.464	49503.082	29037
623	0.453908	-78.268578	804045.402	50227.93	29036
624	0.453617	-78.267486	804167.061	50195.774	29036
625	0.453496	-78.272684	803588.016	50182.166	29036
626	0.447022	-78.273741	803470.536	49465.734	29036
627	0.447224	-78.273284	803521.437	49488.105	29036
628	0.454261	-78.269952	803892.325	50266.933	29035
629	0.453733	-78.267045	804216.183	50208.629	29035
630	0.448777	-78.272251	803636.448	49659.997	29035
631	0.450154	-78.273424	803505.72	49812.322	29035

Station ID	Latitude	Longitude	UTM_X	UTM_Y	Magnetic_data (nT)
632	0.453328	-78.272417	803617.767	50163.587	29034
633	0.452222	-78.272778	803577.598	50041.186	29034
634	0.446557	-78.270494	803832.266	49414.413	29034
635	0.447726	-78.268723	804029.506	49543.844	29034
636	0.444535	-78.270457	803836.471	49190.667	29034
637	0.456991	-78.270885	803788.276	50568.986	29033
638	0.449357	-78.27305	803547.416	49724.145	29033
639	0.447345	-78.269408	803953.213	49501.655	29033
640	0.452692	-78.267357	804181.47	50093.422	29033
641	0.453996	-78.269686	803921.968	50237.621	29032
642	0.454081	-78.269456	803947.586	50247.036	29032
643	0.453964	-78.269261	803969.314	50234.098	29032
644	0.447075	-78.274245	803414.389	49471.578	29032
645	0.444501	-78.272628	803594.626	49186.815	29032
646	0.44844	-78.267258	804192.676	49622.914	29032
647	0.453746	-78.267246	804193.792	50210.059	29031
648	0.452515	-78.269992	803887.942	50073.725	29031
649	0.447468	-78.271605	803708.465	49515.175	29031
650	0.448205	-78.269876	803901.043	49596.801	29031
651	0.444392	-78.272367	803623.706	49174.765	29031
652	0.447056	-78.271958	803669.158	49469.57	29031
653	0.44783	-78.268729	804028.833	49555.352	29031
654	0.448023	-78.268837	804016.794	49576.704	29031
655	0.45116	-78.27344	803503.896	49923.642	29030
656	0.458958	-78.271137	803760.12	50786.637	29030
657	0.45425	-78.269683	803922.292	50265.728	29030
658	0.452792	-78.27314	803537.248	50104.245	29030
659	0.447192	-78.271231	803750.14	49484.649	29030
660	0.44711	-78.272715	803584.827	49475.514	29030
661	0.452531	-78.276036	803214.65	50075.243	29029
662	0.453565	-78.266821	804241.144	50190.048	29029
663	0.453235	-78.272752	803580.452	50153.282	29029
664	0.451374	-78.271701	803697.609	49947.395	29029
665	0.449471	-78.273489	803498.508	49736.741	29029
666	0.446945	-78.27025	803859.432	49457.358	29029
667	0.445397	-78.272338	803626.895	49285.975	29029
668	0.447646	-78.269709	803919.67	49534.951	29029
669	0.445621	-78.270086	803877.756	49310.855	29029
670	0.451373	-78.273833	803460.108	49947.195	29028
671	0.456338	-78.271508	803718.902	50496.701	29028
672	0.455762	-78.269202	803975.811	50433.061	29028
673	0.448564	-78.272421	803617.519	49636.421	29028
674	0.446251	-78.270932	803783.486	49380.534	29028
675	0.451517	-78.268247	804082.374	49963.363	29028
676	0.453583	-78.269068	803990.83	50191.946	29027
677	0.452952	-78.272229	803631.93	50121.986	29027
678	0.448863	-78.272833	803571.61	49669.49	29027
679	0.449142	-78.272942	803559.456	49700.358	29027
680	0.447431	-78.273543	803492.576	49511	29027
681	0.448111	-78.270378	803845.125	49586.378	29027
682	0.444266	-78.272725	803581.045	49160.806	29027
683	0.447001	-78.272292	803631.954	49463.47	29027
684	0.44697	-78.271694	803698.571	49460.064	29027

Station ID	Latitude	Longitude	UTM_X	UTM_Y	Magnetic_data (nT)
685	0.446967	-78.271359	803735.89	49459.746	29027
686	0.453489	-78.266994	804221.875	50181.631	29026
687	0.447472	-78.271869	803679.056	49515.607	29026
688	0.448884	-78.273076	803544.54	49671.803	29026
689	0.448541	-78.270664	803813.247	49633.948	29026
690	0.453762	-78.269281	803967.094	50211.744	29025
691	0.448898	-78.265957	804337.587	49673.649	29025
692	0.450351	-78.266325	804296.532	49834.418	29025
693	0.451386	-78.273583	803487.957	49948.644	29024
694	0.452663	-78.270951	803781.105	50090.062	29024
695	0.449252	-78.273147	803536.615	49712.522	29024
696	0.44754	-78.268796	804021.381	49523.259	29024
697	0.45193	-78.274381	803399.038	50008.808	29023
698	0.452361	-78.274512	803384.427	50056.495	29023
699	0.454295	-78.268132	804095.069	50270.772	29023
700	0.44953	-78.272984	803554.761	49743.291	29023
701	0.448623	-78.271085	803766.345	49643.005	29023
702	0.447173	-78.274026	803438.782	49482.431	29023
703	0.447931	-78.267553	804159.834	49566.577	29023
704	0.45286	-78.272676	803588.934	50111.789	29022
705	0.447702	-78.271702	803697.65	49541.065	29022
706	0.448408	-78.270905	803786.405	49619.221	29022
707	0.448744	-78.270095	803876.624	49656.435	29022
708	0.450585	-78.272515	803606.964	49860.053	29021
709	0.451441	-78.272294	803631.547	49954.784	29021
710	0.448478	-78.271318	803740.395	49626.95	29021
711	0.447912	-78.271534	803716.356	49564.309	29021
712	0.447344	-78.270876	803789.68	49501.484	29021
713	0.445305	-78.27195	803670.121	49275.811	29021
714	0.444975	-78.27041	803841.689	49239.358	29021
715	0.458322	-78.269799	803909.198	50716.316	29020
716	0.452718	-78.272885	803565.658	50096.067	29020
717	0.448228	-78.268727	804029.039	49599.393	29020
718	0.452465	-78.266624	804263.135	50068.334	29020
719	0.453534	-78.269764	803913.298	50186.494	29019
720	0.453816	-78.269466	803946.483	50217.712	29019
721	0.451779	-78.273241	803526.039	49992.146	29019
722	0.447993	-78.272493	803609.522	49573.233	29019
723	0.447532	-78.273767	803467.619	49522.167	29019
724	0.447864	-78.27022	803862.736	49559.052	29019
725	0.447083	-78.270986	803777.437	49472.598	29019
726	0.446523	-78.271925	803672.856	49410.592	29019
727	0.445334	-78.271466	803724.037	49279.04	29019
728	0.451584	-78.267174	804201.903	49970.822	29019
729	0.453897	-78.267854	804126.055	50226.743	29018
730	0.44844	-78.273611	803484.96	49622.65	29018
731	0.447547	-78.27233	803519.641	49523.847	29018
732	0.447025	-78.269017	803996.783	49466.262	29018
733	0.447581	-78.268506	804053.685	49527.808	29018
734	0.44865	-78.267431	804173.395	49646.145	29018
735	0.449197	-78.272183	803644.005	49706.476	29017
736	0.448909	-78.273381	803510.562	49674.557	29017
737	0.446913	-78.27399	803442.803	49453.662	29017

Station ID	Latitude	Longitude	UTM_X	UTM_Y	Magnetic_data (nT)
738	0.445108	-78.271062	803769.052	49254.048	29017
739	0.450235	-78.273158	803535.349	49821.296	29016
740	0.454265	-78.268658	804036.475	50267.431	29016
741	0.452571	-78.2739	803452.594	50079.758	29016
742	0.452368	-78.270255	803858.65	50057.448	29016
743	0.448565	-78.272105	803652.721	49636.544	29016
744	0.449646	-78.273295	803520.112	49756.114	29016
745	0.448108	-78.269271	803968.443	49586.092	29016
746	0.451622	-78.268401	804065.215	49974.976	29016
747	0.452215	-78.273226	803527.692	50040.393	29015
748	0.45074	-78.271113	803763.138	49877.263	29015
749	0.450831	-78.271553	803714.119	49887.315	29015
750	0.448299	-78.270583	803822.28	49607.173	29015
751	0.448167	-78.270822	803795.661	49592.556	29015
752	0.447695	-78.270434	803838.904	49540.343	29015
753	0.445261	-78.2698	803909.63	49271.031	29015
754	0.453554	-78.269502	803942.484	50188.718	29014
755	0.453652	-78.268564	804046.972	50199.602	29014
756	0.451508	-78.272481	803610.713	49962.19	29014
757	0.448336	-78.272578	803600.039	49611.184	29014
758	0.447629	-78.274135	803426.62	49532.886	29014
759	0.4475	-78.270869	803790.453	49518.747	29014
760	0.448001	-78.269976	803889.912	49574.223	29014
761	0.447256	-78.268033	804106.39	49491.864	29014
762	0.444427	-78.270853	803792.362	49178.7	29014
763	0.453378	-78.270794	803798.564	50169.188	29013
764	0.453052	-78.271902	803675.148	50133.068	29013
765	0.45083	-78.272609	803596.482	49887.16	29013
766	0.447298	-78.274296	803408.699	49496.252	29013
767	0.444676	-78.270858	803791.795	49206.253	29013
768	0.444805	-78.27062	803818.302	49220.538	29013
769	0.449322	-78.272754	803580.392	49720.284	29012
770	0.44805	-78.272117	803651.405	49579.556	29012
771	0.445451	-78.273187	803532.315	49291.916	29012
772	0.447433	-78.269951	803892.72	49511.371	29012
773	0.450479	-78.266875	804235.257	49848.559	29012
774	0.454095	-78.268489	804055.308	50248.626	29011
775	0.453968	-78.267643	804149.557	50234.608	29011
776	0.447829	-78.27227	803634.37	49555.094	29011
777	0.448702	-78.273463	803501.436	49651.648	29011
778	0.447283	-78.272179	803644.53	49494.68	29011
779	0.447146	-78.269236	803972.382	49479.642	29011
780	0.445056	-78.271927	803672.694	49248.259	29011
781	0.448273	-78.271957	803669.22	49604.239	29010
782	0.447907	-78.271886	803677.144	49563.742	29010
783	0.449309	-78.273699	803475.121	49718.806	29010
784	0.448202	-78.273677	803477.617	49596.311	29010
785	0.446012	-78.270775	803800.986	49354.094	29010
786	0.450546	-78.267144	804205.288	49855.962	29010
787	0.452537	-78.274338	803403.803	50075.978	29009
788	0.45618	-78.270688	803810.255	50479.252	29009
789	0.456128	-78.271416	803729.159	50473.467	29009
790	0.445891	-78.264928	804452.341	49340.946	29009

Station ID	Latitude	Longitude	UTM_X	UTM_Y	Magnetic_data (nT)
791	0.448366	-78.267553	804159.816	49614.713	29009
792	0.454137	-78.267924	804118.247	50253.297	29008
793	0.452019	-78.271678	803700.145	50018.769	29008
794	0.448816	-78.273615	803484.499	49664.256	29008
795	0.445021	-78.270788	803799.578	49244.433	29008
796	0.451158	-78.266221	804308.084	49923.722	29008
797	0.452419	-78.27283	803571.797	50062.983	29007
798	0.449029	-78.272394	803620.507	49687.877	29007
799	0.448584	-78.272708	803585.547	49638.622	29007
800	0.448109	-78.272714	803584.898	49586.06	29007
801	0.44877	-78.270852	803792.295	49659.281	29007
802	0.448434	-78.273869	803456.219	49621.975	29007
803	0.447455	-78.270398	803842.924	49513.787	29007
804	0.445373	-78.270575	803823.292	49283.392	29007
805	0.450921	-78.273555	803491.095	49897.19	29006
806	0.453045	-78.274021	803439.095	50132.204	29006
807	0.452382	-78.273079	803544.06	50058.879	29006
808	0.448257	-78.272945	803559.159	49602.427	29006
809	0.443667	-78.270896	803787.603	49094.599	29006
810	0.445076	-78.269649	803926.459	49250.566	29006
811	0.451436	-78.266024	804330.018	49954.493	29006
812	0.44977	-78.273118	803539.824	49769.843	29005
813	0.451126	-78.272985	803554.584	49919.898	29005
814	0.452586	-78.273545	803492.14	50081.433	29005
815	0.450606	-78.271342	803737.633	49862.426	29005
816	0.448892	-78.271676	803700.497	49672.747	29005
817	0.448614	-78.273016	803551.235	49641.929	29005
818	0.448468	-78.270061	803880.423	49625.896	29005
819	0.448051	-78.269122	803985.044	49579.791	29005
820	0.445313	-78.271029	803772.719	49276.734	29005
821	0.45008	-78.265996	804333.193	49804.444	29005
822	0.450834	-78.272815	803573.534	49887.594	29004
823	0.450373	-78.271248	803748.114	49836.647	29004
824	0.449035	-78.270409	803841.633	49688.623	29004
825	0.446459	-78.270078	803878.612	49403.586	29004
826	0.445481	-78.270316	803852.14	49295.354	29004
827	0.444473	-78.269371	803957.453	49183.851	29004
828	0.444344	-78.268738	804027.974	49169.603	29004
829	0.452031	-78.274095	803430.894	50019.996	29003
830	0.450389	-78.270752	803803.367	49838.438	29003
831	0.450757	-78.272244	803637.146	49879.097	29003
832	0.44812	-78.272422	803617.426	49587.289	29003
833	0.44778	-78.273896	803453.238	49549.605	29003
834	0.447343	-78.270631	803816.973	49501.383	29003
835	0.446586	-78.269921	803896.097	49417.646	29003
836	0.446735	-78.271913	803674.185	49434.051	29003
837	0.448217	-78.268297	804076.941	49598.194	29003
838	0.444816	-78.270163	803869.211	49221.774	29003
839	0.444128	-78.27036	803847.294	49145.634	29003
840	0.45283	-78.271984	803666.023	50108.499	29002
841	0.450981	-78.271321	803739.957	49903.923	29002
842	0.450837	-78.272	803664.323	49887.96	29002
843	0.448516	-78.271775	803689.484	49631.136	29002

Station ID	Latitude	Longitude	UTM_X	UTM_Y	Magnetic_data (nT)
1268	0.445942	-78.265956	804337.82	49346.547	28934
1269	0.446258	-78.265455	804393.618	49381.535	28934
1270	0.451943	-78.270241	803860.228	50010.419	28934
1271	0.451805	-78.271901	803675.312	49995.079	28934
1272	0.446336	-78.267872	804124.363	49390.067	28934
1273	0.448835	-78.267852	804126.488	49666.599	28934
1274	0.445109	-78.268427	804062.587	49254.268	28934
1275	0.446078	-78.268895	804010.413	49361.475	28934
1276	0.445297	-78.269268	803968.893	49275.036	28934
1277	0.447125	-78.266623	804263.468	49477.426	28933
1278	0.44659	-78.266533	804273.516	49418.229	28933
1279	0.446786	-78.265214	804420.444	49439.972	28933
1280	0.449571	-78.271491	803721.078	49747.89	28933
1281	0.450718	-78.269659	803925.112	49874.89	28933
1282	0.452173	-78.26733	804184.5	50035.993	28933
1283	0.446262	-78.269105	803987.012	49381.827	28933
1284	0.450194	-78.267335	804184.025	49817.003	28933
1285	0.450377	-78.26971	803919.445	49837.154	28932
1286	0.449295	-78.268981	804000.7	49717.454	28932
1287	0.451329	-78.268945	804004.626	49942.531	28932
1288	0.451766	-78.267981	804111.996	49990.928	28932
1289	0.447904	-78.266774	804246.615	49563.622	28931
1290	0.45142	-78.269657	803925.306	49952.571	28931
1291	0.449493	-78.270217	803863.003	49739.312	28931
1292	0.446728	-78.267909	804120.225	49433.442	28931
1293	0.446392	-78.269492	803943.895	49396.196	28931
1294	0.447265	-78.264758	804471.222	49492.996	28931
1295	0.459178	-78.270568	803823.497	50811.005	28930
1296	0.446803	-78.265505	804388.026	49441.841	28930
1297	0.448906	-78.268915	804008.068	49674.411	28930
1298	0.447303	-78.266451	804282.621	49497.131	28929
1299	0.452793	-78.266519	804274.818	50104.634	28929
1300	0.451661	-78.266913	804230.975	49979.354	28929
1301	0.447841	-78.266263	804303.542	49556.672	28928
1302	0.445231	-78.26623	804307.326	49267.859	28928
1303	0.446535	-78.265472	804391.713	49412.187	28928
1304	0.451246	-78.270833	803794.309	49933.267	28928
1305	0.451933	-78.269425	803951.129	50009.347	28928
1306	0.452375	-78.266341	804294.665	50058.387	28928
1307	0.449587	-78.267672	804146.509	49749.82	28928
1308	0.44419	-78.268518	804052.488	49152.571	28928
1309	0.44761	-78.266757	804248.521	49531.089	28927
1310	0.445699	-78.2651	804433.188	49319.693	28927
1311	0.44639	-78.268696	804032.568	49396.008	28927
1312	0.449041	-78.266835	804239.772	49689.436	28927
1313	0.450853	-78.266385	804289.827	49889.965	28927
1314	0.450687	-78.265626	804374.386	49871.628	28927
1315	0.45064	-78.269299	803965.219	49866.273	28926
1316	0.451421	-78.269499	803942.907	49952.688	28926
1317	0.45065	-78.268424	804062.693	49867.417	28926
1318	0.451243	-78.26836	804069.798	49933.039	28926
1319	0.449917	-78.267085	804211.886	49786.361	28926
1320	0.449071	-78.267939	804116.787	49692.71	28926

Station ID	Latitude	Longitude	UTM_X	UTM_Y	Magnetic_data (nT)
1321	0.445629	-78.268594	804043.962	49311.802	28926
1322	0.445965	-78.265168	804425.602	49349.125	28925
1323	0.451782	-78.271575	803711.629	49992.548	28925
1324	0.445777	-78.269347	803960.073	49328.148	28925
1325	0.445623	-78.265593	804378.271	49311.263	28924
1326	0.44616	-78.265686	804367.889	49370.681	28924
1327	0.451027	-78.270669	803812.587	49909.04	28924
1328	0.449095	-78.271806	803686.007	49695.205	28924
1329	0.451725	-78.269265	803968.962	49986.337	28924
1330	0.451463	-78.269197	803976.548	49957.348	28924
1331	0.446681	-78.268623	804040.689	49428.212	28924
1332	0.444543	-78.271475	803723.067	49191.51	28924
1333	0.446791	-78.267131	804206.891	49440.446	28923
1334	0.446664	-78.26688	804234.858	49426.403	28923
1335	0.448271	-78.265005	804443.665	49604.307	28923
1336	0.446616	-78.266204	804310.165	49421.12	28923
1337	0.44628	-78.266468	804280.77	49383.928	28923
1338	0.449334	-78.270983	803777.678	49721.686	28923
1339	0.449057	-78.269449	803948.575	49691.098	28923
1340	0.452183	-78.267608	804153.531	50037.087	28923
1341	0.449387	-78.266896	804232.963	49727.721	28923
1342	0.450931	-78.265894	804344.52	49898.617	28923
1343	0.445881	-78.265668	804369.906	49339.809	28922
1344	0.446	-78.265396	804400.201	49352.988	28922
1345	0.447249	-78.273035	803549.174	49490.882	28922
1346	0.450684	-78.269046	803993.401	49871.153	28922
1347	0.450075	-78.268597	804043.445	49803.782	28922
1348	0.449677	-78.26824	804083.231	49759.755	28922
1349	0.445358	-78.267112	804209.067	49281.876	28922
1350	0.44665	-78.267468	804169.356	49424.829	28921
1351	0.447489	-78.266291	804300.438	49517.719	28921
1352	0.446493	-78.266452	804282.543	49407.498	28921
1353	0.450385	-78.269197	803976.593	49838.06	28921
1354	0.4464	-78.266716	804253.138	49397.196	28920
1355	0.446495	-78.265999	804333.007	49407.739	28920
1356	0.449169	-78.268932	804006.164	49703.513	28919
1357	0.448401	-78.26822	804085.511	49618.558	28919
1358	0.450233	-78.268037	804105.821	49821.289	28919
1359	0.445451	-78.265859	804348.646	49292.219	28918
1360	0.449645	-78.269995	803887.727	49756.141	28918
1361	0.45109	-78.268039	804105.563	49916.122	28918
1362	0.446827	-78.266394	804288.991	49444.46	28917
1363	0.447056	-78.265886	804345.572	49469.822	28917
1364	0.449753	-78.269278	803967.595	49768.122	28917
1365	0.450433	-78.268931	804006.223	49843.383	28917
1366	0.446108	-78.269322	803962.844	49364.777	28917
1367	0.444853	-78.266577	804268.686	49226.016	28917
1368	0.447423	-78.266959	804226.026	49510.388	28916
1369	0.448344	-78.265277	804413.361	49612.373	28916
1370	0.446159	-78.264999	804444.42	49370.599	28916
1371	0.450736	-78.267437	804172.64	49876.974	28916
1372	0.452205	-78.266564	804269.83	50039.566	28916
1373	0.444889	-78.267263	804192.265	49229.971	28916

Station ID	Latitude	Longitude	UTM_X	UTM_Y	Magnetic_data (nT)
1374	0.44643	-78.26571	804365.204	49400.558	28915
1375	0.446512	-78.265207	804421.235	49409.653	28915
1376	0.446984	-78.265036	804440.264	49461.89	28915
1377	0.449339	-78.270552	803825.691	49722.257	28915
1378	0.446255	-78.268142	804094.289	49381.092	28915
1379	0.451955	-78.267611	804153.206	50011.858	28915
1380	0.449748	-78.267784	804134.026	49767.631	28915
1381	0.443993	-78.268691	804033.224	49130.764	28915
1382	0.446145	-78.264504	804499.563	49369.071	28915
1383	0.45199	-78.266614	804264.269	50015.772	28915
1384	0.445848	-78.268749	804026.687	49336.03	28914
1385	0.445229	-78.267259	804192.696	49267.595	28914
1386	0.4458	-78.267799	804132.517	49330.758	28914
1387	0.449664	-78.265607	804376.545	49758.427	28914
1388	0.4479	-78.265377	804402.239	49563.237	28913
1389	0.452319	-78.268428	804062.178	50052.102	28913
1390	0.449785	-78.266856	804237.402	49771.764	28913
1391	0.444566	-78.268147	804093.801	49194.193	28913
1392	0.446965	-78.266941	804228.05	49459.708	28912
1393	0.444951	-78.266195	804311.236	49236.876	28912
1394	0.45077	-78.26777	804135.543	49880.723	28912
1395	0.4462	-78.264756	804471.488	49375.146	28912
1396	0.446701	-78.265729	804363.076	49430.545	28911
1397	0.445729	-78.26534	804406.451	49323.003	28911
1398	0.446703	-78.265024	804441.613	49430.796	28911
1399	0.450077	-78.267546	804160.525	49804.047	28911
1400	0.449744	-78.268009	804108.961	49767.179	28911
1401	0.445255	-78.268972	804001.869	49270.401	28911
1402	0.445507	-78.267331	804184.664	49298.354	28911
1403	0.446456	-78.264741	804473.149	49403.475	28911
1404	0.448174	-78.264952	804449.573	49593.575	28910
1405	0.45079	-78.270033	803883.446	49882.841	28910
1406	0.450872	-78.267071	804213.406	49892.039	28910
1407	0.446594	-78.268373	804068.542	49418.595	28910
1408	0.450227	-78.267778	804134.451	49820.636	28910
1409	0.44521	-78.267588	804156.047	49265.479	28910
1410	0.447053	-78.265237	804417.87	49469.517	28909
1411	0.449244	-78.267384	804178.606	49711.877	28909
1412	0.44772	-78.266947	804227.35	49543.254	28908
1413	0.448104	-78.265518	804386.524	49585.806	28908
1414	0.449189	-78.269213	803974.86	49705.714	28908
1415	0.449213	-78.266341	804294.796	49708.49	28908
1416	0.44523	-78.269422	803951.74	49267.616	28908
1417	0.449844	-78.265527	804385.449	49778.348	28908
1418	0.451736	-78.27131	803741.151	49987.469	28907
1419	0.44922	-78.268495	804054.843	49709.175	28907
1420	0.445056	-78.268795	804021.595	49248.388	28907
1421	0.449741	-78.265295	804411.298	49766.96	28907
1422	0.448066	-78.266426	804285.375	49581.563	28906
1423	0.447188	-78.267044	804216.567	49484.38	28906
1424	0.447348	-78.266141	804317.153	49502.123	28906
1425	0.449627	-78.270544	803826.57	49754.126	28906
1426	0.44874	-78.268697	804032.36	49656.051	28906

Station ID	Latitude	Longitude	UTM_X	UTM_Y	Magnetic_data (nT)
1427	0.448996	-78.268375	804068.22	49684.392	28906
1428	0.451661	-78.269027	803995.477	49979.265	28906
1429	0.449333	-78.265374	804402.514	49721.809	28906
1430	0.444518	-78.268525	804051.694	49188.866	28906
1431	0.445602	-78.267615	804153.023	49308.855	28906
1432	0.446571	-78.267788	804133.711	49416.074	28905
1433	0.449258	-78.26673	804251.46	49713.453	28905
1434	0.44383	-78.268969	804002.261	49112.716	28905
1435	0.44517	-78.265979	804335.29	49261.119	28905
1436	0.446877	-78.266729	804251.67	49449.979	28904
1437	0.44519	-78.265781	804357.346	49263.34	28904
1438	0.44982	-78.269081	803989.538	49775.544	28904
1439	0.446473	-78.266972	804224.617	49405.264	28903
1440	0.449244	-78.26582	804352.834	49711.942	28903
1441	0.444674	-78.268735	804028.294	49206.119	28903
1442	0.443833	-78.268476	804057.181	49113.068	28903
1443	0.443364	-78.269166	803980.335	49061.141	28903
1444	0.452289	-78.268173	804090.586	50048.793	28902
1445	0.445266	-78.268645	804038.296	49271.632	28902
1446	0.445946	-78.264666	804481.525	49347.043	28902
1447	0.447517	-78.267119	804208.198	49520.783	28901
1448	0.452035	-78.267884	804122.791	50020.699	28901
1449	0.445059	-78.267466	804169.644	49248.775	28901
1450	0.449907	-78.269536	803938.848	49785.152	28900
1451	0.449123	-78.268211	804086.484	49698.453	28900
1452	0.448128	-78.266678	804257.3	49588.413	28899
1453	0.447854	-78.26659	804267.114	49558.097	28899
1454	0.452077	-78.266804	804243.1	50025.392	28899
1455	0.448216	-78.266215	804308.874	49598.17	28899
1456	0.449732	-78.266569	804269.376	49765.911	28899
1457	0.445365	-78.265521	804386.302	49282.716	28898
1458	0.449286	-78.267107	804209.462	49716.536	28898
1459	0.451853	-78.266834	804239.767	50000.603	28898
1460	0.447007	-78.267228	804196.077	49464.344	28897
1461	0.44726	-78.26506	804437.579	49492.43	28897
1462	0.451502	-78.268667	804035.587	49961.686	28897
1463	0.451003	-78.267386	804178.31	49906.522	28897
1464	0.44772	-78.265722	804363.814	49543.305	28897
1465	0.451391	-78.271188	803754.756	49949.297	28896
1466	0.450547	-78.269498	803943.055	49855.974	28896
1467	0.448681	-78.273733	803471.359	49649.313	28896
1468	0.444917	-78.266322	804297.09	49233.109	28896
1469	0.450624	-78.266394	804288.834	49864.624	28896
1470	0.446499	-78.268107	804098.178	49408.094	28895
1471	0.445044	-78.269141	803983.051	49247.046	28895
1472	0.451829	-78.267123	804207.574	49997.935	28895
1473	0.45169	-78.266633	804262.165	49982.575	28895
1474	0.44671	-78.267715	804141.838	49431.459	28894
1475	0.451524	-78.268779	804023.11	49964.116	28894
1476	0.444733	-78.26836	804070.066	49212.664	28894
1477	0.443475	-78.26892	804007.735	49073.434	28894
1478	0.44641	-78.264474	804502.894	49398.396	28894
1479	0.448023	-78.265098	804433.315	49576.86	28893

Station ID	Latitude	Longitude	UTM_X	UTM_Y	Magnetic_data (nT)
1480	0.449778	-78.26607	804324.962	49771.022	28893
1481	0.449661	-78.266206	804309.817	49758.07	28893
1482	0.449268	-78.26799	804111.097	49714.507	28893
1483	0.449143	-78.267672	804146.527	49700.688	28893
1484	0.445269	-78.266957	804226.337	49272.034	28893
1485	0.446833	-78.266038	804328.649	49445.139	28892
1486	0.44737	-78.272811	803574.123	49504.281	28892
1487	0.450924	-78.267538	804161.381	49897.774	28892
1488	0.449499	-78.266363	804292.334	49740.137	28892
1489	0.445179	-78.268063	804103.134	49262.029	28892
1490	0.443659	-78.268251	804082.253	49093.823	28891
1491	0.451831	-78.266392	804289.006	49998.187	28891
1492	0.449616	-78.26583	804351.705	49753.106	28890
1493	0.447445	-78.265663	804370.398	49512.877	28889
1494	0.450064	-78.267728	804140.251	49802.601	28889
1495	0.449749	-78.267188	804200.419	49767.766	28889
1496	0.45246	-78.268605	804042.454	50067.698	28888
1497	0.444462	-78.267568	804158.306	49182.708	28888
1498	0.444348	-78.268312	804075.429	49170.063	28888
1499	0.449639	-78.26741	804175.693	49755.585	28887
1500	0.444794	-78.267998	804110.39	49219.429	28887
1501	0.450537	-78.269864	803902.283	49854.852	28886
1502	0.447595	-78.267203	804198.837	49529.411	28885
1503	0.447231	-78.266009	804331.863	49489.182	28885
1504	0.443661	-78.268755	804026.108	49094.023	28884
1505	0.450893	-78.26726	804192.351	49894.355	28883
1506	0.449361	-78.266549	804271.619	49724.858	28883
1507	0.447489	-78.265924	804341.321	49517.735	28882
1508	0.451865	-78.269928	803895.099	50001.801	28882
1509	0.450842	-78.266098	804321.799	49888.76	28882
1510	0.451924	-78.266205	804309.834	50008.486	28882
1511	0.447661	-78.2655	804388.547	49536.785	28881
1512	0.445891	-78.264401	804511.048	49340.968	28880
1513	0.459111	-78.27107	803767.578	50803.57	28879
1514	0.449014	-78.268196	804088.16	49686.392	28879
1515	0.451911	-78.267403	804176.379	50006.997	28879
1516	0.447829	-78.267062	804214.535	49555.311	28878
1517	0.449453	-78.265614	804375.774	49735.078	28878
1518	0.444655	-78.2717	803697.998	49203.895	28878
1519	0.443509	-78.268537	804050.399	49077.213	28878
1520	0.447355	-78.267306	804187.373	49502.849	28877
1521	0.451245	-78.267916	804119.259	49933.279	28877
1522	0.444019	-78.268311	804075.554	49133.657	28877
1523	0.451612	-78.271745	803692.698	49973.729	28875
1524	0.445944	-78.268154	804092.965	49346.677	28875
1525	0.451549	-78.271513	803718.545	49966.768	28873
1526	0.451041	-78.266209	804309.425	49910.776	28873
1527	0.446951	-78.267636	804150.628	49458.13	28872
1528	0.450784	-78.266897	804232.793	49882.308	28872
1529	0.444891	-78.268577	804045.886	49230.138	28872
1530	0.451967	-78.266038	804328.436	50013.251	28872
1531	0.451243	-78.271399	803731.257	49932.912	28871
1532	0.444007	-78.269164	803980.531	49132.294	28871

Station ID	Latitude	Longitude	UTM_X	UTM_Y	Magnetic_data (nT)
1533	0.4495	-78.265192	804422.782	49740.296	28871
1534	0.447541	-78.272587	803599.069	49523.212	28870
1535	0.449527	-78.266776	804246.325	49743.218	28870
1536	0.447929	-78.265682	804368.262	49566.434	28869
1537	0.451082	-78.267776	804134.861	49915.247	28869
1538	0.450824	-78.266651	804260.196	49886.745	28869
1539	0.452079	-78.26708	804212.353	50025.601	28868
1540	0.444632	-78.267778	804134.905	49201.511	28868
1541	0.451061	-78.267109	804209.165	49912.951	28867
1542	0.444956	-78.268208	804086.99	49237.346	28867
1543	0.445787	-78.268373	804068.575	49329.295	28867
1544	0.450318	-78.269385	803955.652	49830.638	28866
1545	0.449475	-78.267223	804196.532	49737.445	28866
1546	0.444402	-78.267934	804117.536	49176.054	28866
1547	0.443526	-78.26938	803956.489	49079.059	28866
1548	0.452706	-78.276226	803193.477	50094.6	28864
1549	0.44935	-78.267748	804138.052	49723.591	28864
1550	0.444851	-78.26764	804150.269	49225.751	28864
1551	0.445007	-78.267851	804126.757	49243.005	28864
1552	0.446362	-78.269892	803899.337	49392.86	28864
1553	0.448963	-78.271055	803769.673	49680.629	28861
1554	0.443869	-78.268089	804100.291	49117.067	28861
1555	0.449648	-78.265042	804439.486	49756.68	28857
1556	0.449927	-78.268265	804080.435	49787.418	28856
1557	0.450595	-78.266879	804234.806	49861.395	28853
1558	0.443306	-78.268707	804031.469	49054.742	28850
1559	0.4449	-78.265939	804339.757	49231.243	28850
1560	0.445435	-78.267829	804129.19	49290.367	28848
1561	0.445097	-78.265471	804391.883	49253.062	28847
1562	0.449562	-78.266976	804224.044	49747.082	28843
1563	0.447009	-78.266174	804313.491	49464.609	28841
1564	0.449426	-78.266008	804331.883	49732.074	28841
1565	0.444193	-78.268093	804099.832	49152.92	28840
1566	0.448117	-78.265212	804420.611	49587.257	28839
1567	0.444679	-78.267424	804174.338	49206.727	28835
1568	0.444296	-78.267351	804182.486	49164.348	28835
1569	0.449393	-78.267488	804167.014	49728.36	28834
1570	0.450594	-78.268706	804031.281	49861.208	28833
1571	0.451326	-78.268134	804094.97	49942.233	28833
1572	0.45919	-78.270845	803792.639	50812.322	28830
1573	0.445641	-78.264821	804464.271	49313.286	28830
1574	0.449974	-78.267977	804112.516	49792.631	28829
1575	0.4457	-78.268972	804001.851	49319.643	28829
1576	0.444926	-78.265694	804367.048	49234.13	28825
1577	0.444245	-78.267719	804141.493	49158.69	28823
1578	0.444028	-78.267875	804124.124	49134.671	28821
1579	0.445445	-78.265002	804444.115	49291.59	28820
1580	0.448046	-78.268506	804053.666	49579.263	28806
1581	0.448872	-78.268442	804060.761	49670.668	28805
1582	0.445692	-78.264561	804493.232	49318.941	28803
1583	0.445501	-78.264715	804476.085	49297.799	28802
1584	0.44688	-78.269374	803957.02	49450.202	28796
1585	0.445681	-78.267502	804165.608	49317.602	28795

Station ID	Latitude	Longitude	UTM_X	UTM_Y	Magnetic_data (nT)
1586	0.445204	-78.265159	804426.636	49264.915	28794
1587	0.448309	-78.264878	804457.811	49608.517	28790
1588	0.444071	-78.267507	804165.117	49139.444	28789
1589	0.444718	-78.265635	804373.629	49211.116	28780
1590	0.443912	-78.26729	804189.297	49121.859	28780
1591	0.444134	-78.267138	804206.221	49146.431	28779
1592	0.443956	-78.26693	804229.399	49126.742	28759
1593	0.452961	-78.276625	803149.018	50122.8	28747
1594	0.445184	-78.264928	804452.37	49262.712	28735
1595	0.444821	-78.265385	804401.475	49222.524	28734
1596	0.449379	-78.272517	803606.791	49726.601	28723
1597	0.444962	-78.265084	804435	49238.139	28722
1598	0.443736	-78.26708	804212.698	49102.392	28715
1599	0.452762	-78.276703	803140.337	50100.776	28708
1600	0.448385	-78.269225	803973.556	49616.746	28672
1601	0.443791	-78.266718	804253.022	49108.493	28665
1602	0.44357	-78.266872	804235.876	49084.031	28643
1603	0.448801	-78.269367	803957.72	49662.773	28631
1604	0.450229	-78.267495	804166.2	49820.869	28627
1605	0.443407	-78.266663	804259.165	49066.003	28626
1606	0.451671	-78.272157	803646.799	49980.241	28620
1607	0.451649	-78.272161	803646.355	49977.806	28600
1608	0.443623	-78.266507	804276.534	49089.911	28581
1609	0.443246	-78.266454	804282.454	49048.196	28539
1610	0.44346	-78.266293	804300.38	49071.883	28530
1611	0.443296	-78.266075	804324.672	49053.744	28520
1612	0.443082	-78.266242	804306.077	49030.057	28477
1613	0.452944	-78.276441	803169.516	50120.927	28465